

DYNAMIC RESPONSE OF SOFT POROELASTIC BED TO NONLINEAR WATER WAVE—BOUNDARY LAYER CORRECTION APPROACH

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ABSTRACT: When an oscillatory water wave propagates over a soft poroelastic bed, a boundary layer exists within the porous bed and near the homogeneous water/porous bed interface. Owing to the effect of the boundary layer, the conventional evaluation of the second kind of longitudinal wave inside the soft poroelastic bed by one parameter, $\varepsilon_1 = k_0 a$, is very inaccurate so that a boundary layer correction approach for a soft poroelastic bed is proposed to solve the nonlinear water wave problem. Hence a perturbation expansion for the boundary layer correction approach based on two small parameters, ε_1 and $\varepsilon_2 = k_0/k_2$, is proposed and then solved. The solutions carried out to the first three terms are valid for the first kind and the third kind of waves throughout the whole domain. The second kind of wave is solved systematically inside the boundary layer, whereas it disappears outside the boundary layer. The result is compared with the linear wave solution of Huang and Song in order to show the nonlinearity effect. The present study is very helpful to formulate a simplified boundary-value problem in numerical computation for soft poroelastic medium with irregular geometry.

INTRODUCTION

The dynamic action of a propagating water wave on coastal constructions is emphasized during the design work, especially in the analysis of seabed instability. The wave-induced variation in pore pressure and effective stresses has been recognized as a major factor so that it is very important to correctly estimate the stress and strain of the seabed. In general, the seabed is permeable and deforming, and the nonlinear water wave is very likely to happen. The study of the interaction between the seabed and the water wave becomes increasingly more complicated; thus it has been changed from linear water wave to nonlinear water wave and from rigid seabed material to poroelastic bed material.

Putnam (1949) began investigation on a linear water wave interacting with a porous bed. Then, Reid and Kajiura (1957) studied the porous bed problem by considering a linear wave in an inviscid, incompressible, irrotational fluid flow satisfying Darcy's law interacting with a rigid, isotropic porous skeleton. Sleath (1970), Liu (1973), and Moshagen and Torum (1975) improved the studies of Putnam (1949) and Reid and Kajiura (1957), but all the earlier studies focused only on the rigid bed material.

In fact, fluid within porous material interacting with a deforming solid skeleton becomes a more complicated two-phase problem for a realistic analysis. Biot (1962) developed a poroelasticity theory to discuss an elastic wave in a fluid-saturated porous solid. Yamamoto et al. (1978) obtained governing equations of a water wave propagating over a porous bed under the assumption of existing double eigenvalues. These equations are exactly the same as the limiting equation of poroelasticity [see discussions in Huang and Song (1993)]. By neglecting the inertial terms of the poroelasticity theory that is physically reasonable for rigid bed material, Madsen (1978) also applied Biot's theory to investigate the effect of anisotropic permeability in a seabed by a numerical method that is only valid for a rigid porous bed. On the other hand, Mei and Foda (1981) proposed a boundary layer correction to simplify

the analysis; however, their approach was imperfectly assumed to be restricted to a low-frequency wave only and without systematic perturbation analysis. For Biot's equation without simplifications, Huang and Chwang (1990) investigated Biot's oscillatory equation for an acoustic problem and obtained three decoupled Helmholtz equations to represent each of the three kinds of wave. Huang and Song (1993) solved the problem of a linear water wave interacting with a deformable bed by treating the bed as a poroelastic material and obtained some satisfactory results.

As to the nonlinear water wave problem, Mei (1983), Fenton (1985), and Dean and Dalrymple (1991) studied the nonlinearly deep-water wave on an impermeable rigid bed by the Stokes expansion. Chen et al. (1997) also applied the conventional Stokes expansion of the deep-water wave based on $\varepsilon_1 = k_0 a$ to investigate the dynamic response of a permeable bed material. They found that the Stokes expansion used is only valid for hard poroelastic bed material but invalid for a soft one even though the Ursell parameter is small.

Based on the foregoing comments and because the wavelength of the second kind of longitudinal wave (with wave number k_2) inside the porous bed is much shorter than that of the water wave (with wave number k_0) (i.e., $\|k_2\| > \|k_0\|$ for soft porous bed material), a two-parameter expansion based on ε_1 and $\varepsilon_2 = k_0/k_2$ instead of a one-parameter expansion based on ε_1 is proposed to investigate the problem of a nonlinear water wave propagating over a soft poroelastic deforming bed. This approach will be more systematic than the approach of Mei and Foda (1981). Moreover, solving the interaction between water and a soft porous bed is very hard even by a numerical method; thus the present study will propose a basic concept to formulate a simplified boundary-value problem of a soft poroelastic medium for numerical computation in the sequence research, whereas Madsen (1978) only discussed the problem of a rigid porous bed.

FORMULATION

This study on a nonlinear water wave problem is defined in Fig. 1, which indicates that the plane wave propagates over a horizontally infinite thickness and homogeneous poroelastic bed saturated with water. Region 1 is homogeneous water processed as potential flow, and region 2 is a semi-infinite porous medium treated by Biot's theory of poroelasticity (Biot 1962). The symbols η^* and ξ^* in Fig. 1 represent the displacements of waves from the mean free surface ($y = h$) and mean bed interface ($y = 0$), respectively.

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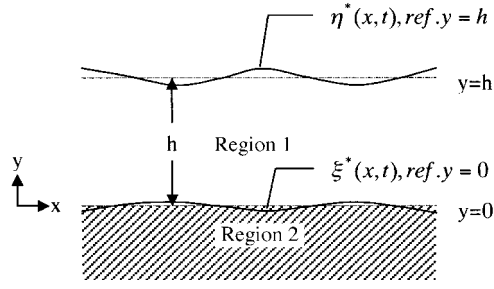


FIG. 1. Definition Sketch

Boundary-Value Problem

Assuming that homogeneous channel flow (region 1 of Fig. 1) is potential flow, the equations of continuity and momentum in terms of velocity potential $\Phi^{*(1)}$ become

$$\nabla^2 \Phi^{*(1)} = 0 \quad (1)$$

$$\rho_0 \frac{\partial \Phi^{*(1)}}{\partial t} + \frac{\rho_0}{2} \left\{ \left[\frac{\partial \Phi^{*(1)}}{\partial x} \right]^2 + \left[\frac{\partial \Phi^{*(1)}}{\partial y} \right]^2 \right\} + P^{*(1)} = 0 \quad (2)$$

where $P^{*(1)}$ = perturbed pressure; and ρ_0 = water density.

Referring to Huang and Chwang (1990), the linear momentum equations of solid skeleton and fluid for the porous bed based on the theory of poroelasticity may be written as

$$\nabla \cdot \boldsymbol{\sigma}^* = \rho_{11} \frac{\partial^2 \mathbf{d}^*}{\partial t^2} + \rho_{12} \frac{\partial^2 \mathbf{D}^*}{\partial t^2} + b \left(\frac{\partial \mathbf{d}^*}{\partial t} - \frac{\partial \mathbf{D}^*}{\partial t} \right) \quad (3)$$

$$\nabla \cdot \mathbf{S}^* = \rho_{12} \frac{\partial^2 \mathbf{d}^*}{\partial t^2} + \rho_{22} \frac{\partial^2 \mathbf{D}^*}{\partial t^2} - b \left(\frac{\partial \mathbf{d}^*}{\partial t} - \frac{\partial \mathbf{D}^*}{\partial t} \right) \quad (4)$$

with

$$\boldsymbol{\sigma}^* = \boldsymbol{\tau}^* - (1 - n_0)P^{*(2)}\mathbf{I} \quad (5)$$

$$\boldsymbol{\tau}^* = 2G\mathbf{e}^* + \lambda(\nabla \cdot \mathbf{d}^*)\mathbf{I} \quad (6)$$

$$\mathbf{e}^* = \frac{1}{2} [\nabla \mathbf{d}^* + (\nabla \mathbf{d}^*)'] \quad (7)$$

$$\mathbf{S}^* = -n_0 P^{*(2)}\mathbf{I} \quad (8)$$

$$\rho_{11} = (1 - n_0)\rho_s + \rho_a \quad (9)$$

$$\rho_{12} = -\rho_a \quad (10)$$

$$\rho_{22} = n_0\rho_0 + \rho_a \quad (11)$$

$$b = \mu n_0^2/k_p \quad (12)$$

where $\boldsymbol{\sigma}^*$ = solid stress tensor, $\boldsymbol{\tau}^*$ = effective stress tensor of solid; \mathbf{S}^* = normal stress tensor of fluid; \mathbf{d}^* and \mathbf{D}^* = solid and fluid displacement vectors, respectively; $P^{*(2)}$ = perturbed pore pressure inside the porous medium; ρ_s = solid density; ρ_a = mass coupling effect (neglected in this study); n_0 = porosity; μ = fluid viscosity; k_p = specific permeability; G and λ = Lamé's constants of elasticity; \mathbf{I} = identity matrix.

Combining continuity equations of solid and fluid with state equation of fluid and after linearization of porosity, one can find

$$\frac{\partial P^{*(2)}}{\partial t} = -\frac{K}{n_0} \left[(1 - n_0)\nabla \cdot \left(\frac{\partial \mathbf{d}^*}{\partial t} \right) + n_0\nabla \cdot \left(\frac{\partial \mathbf{D}^*}{\partial t} \right) \right] \quad (13)$$

for perturbed pore pressure $P^{*(2)}$. In (13), K is the bulk modulus of compressibility of fluid inside the porous bed.

There are three boundaries that require boundary conditions in this study. They are (1) free surface [$y = h + \eta^*(x, t)$]; (2) channel-bed interface [$y = \xi^*(x, t)$]; and (3) deep far field of the porous bed [$\hat{y} \rightarrow -\infty$].

On the free surface, a kinematic boundary condition exists as

$$-\frac{\partial \eta^*}{\partial x} \frac{\partial \Phi^{*(1)}}{\partial x} + \frac{\partial \Phi^{*(1)}}{\partial y} = \frac{\partial \eta^*}{\partial t} \quad (14)$$

and a dynamic boundary condition exists as

$$\frac{\partial \Phi^{*(1)}}{\partial t} + \frac{1}{2} \left\{ \left[\frac{\partial \Phi^{*(1)}}{\partial x} \right]^2 + \left[\frac{\partial \Phi^{*(1)}}{\partial y} \right]^2 \right\} + g\eta^* = 0 \quad (15)$$

On the porous bed interface, the continuity of pressure gives

$$P^{*(1)} = P^{*(2)} \quad (16)$$

and the conservation of fluid flux gives

$$\mathbf{n}_2^* \cdot \left[(1 - n_0) \frac{\partial \mathbf{d}^*}{\partial t} + n_0 \frac{\partial \mathbf{D}^*}{\partial t} \right] = \mathbf{n}_2^* \cdot \nabla \Phi^{*(1)} \quad (17)$$

where \mathbf{n}_2^* = unit normal vector of the porous bed interface. Considering the kinematics of the porous bed interface, one has

$$\frac{\partial \xi^*}{\partial t} = \frac{\partial \mathbf{d}^*}{\partial t} \cdot \left(-\frac{\partial \xi^*}{\partial x}, 1 \right) \quad (18)$$

and (18) will be used to solve ξ^* . The continuity of effective stresses of solid gives

$$\mathbf{n}_2^* \cdot \boldsymbol{\tau}^* = 0 \quad (19)$$

At the far field of porous bed, $y \rightarrow -\infty$, the boundary conditions are vanishing displacement vectors; that is

$$\mathbf{d}^*, \mathbf{D}^* \rightarrow 0 \quad (20)$$

If $|\eta^*|$ and $|\xi^*|$ are much smaller than the relative wavelengths, it is more convenient to shift the boundary conditions at the free surface, $y = h + \eta^*(x, t)$, and the porous bed interface, $y = \xi^*(x, t)$, to $y = h$ and $y = 0$ before solving the boundary-value problem. Conventionally, the Taylor series expansions are applied to the boundary conditions at the free surface [(14) and (15)] and at the porous bed interface [(16), (17), (19), and (18)] by performing, respectively

$$\sum_{m=0}^{\infty} \frac{(\eta^*)^m}{m!} \frac{\partial^m}{\partial y^m} \quad \text{and} \quad \sum_{m=0}^{\infty} \frac{(\xi^*)^m}{m!} \frac{\partial^m}{\partial y^m}$$

As $m = 0$, the boundary-wave problem is linear, but as $m \geq 1$, the problem becomes nonlinear. For nonlinear problems, the above Taylor series expansions at the interface ($y = 0$) are applicable for the first and third kinds of wave but are not applicable for the second kind of wave because there exists a boundary layer that will render errors of the partial derivative in the vertical direction for the second longitudinal wave. That is why Chen et al. (1997) failed to solve the nonlinear problem for soft porous material by only one length scale. To overcome the difficulty, another small parameter, $\varepsilon_2 = k_0/k_2$, other than $\varepsilon_1 = k_0 a$ needs to be proposed. Thus, the vertical coordinate y for the second kind of wave will be enlarged into y' based on this small parameter ε_2 [see (35b)].

Referring to Huang and Song (1993) for the decoupling processes of Biot's equations of poroelasticity (Biot 1962), the governing equations (3) and (4) can be rewritten into three decoupled scalar equations as

$$\nabla^2 \Phi_j^{*(2)} + k_j^2 \Phi_j^{*(2)} = 0, \quad j = 1, 2, 3 \quad (21)$$

Also, the perturbed pressure equation [(13)] gives

$$P^{*(2)} = \frac{K}{n_0} \left[(1 - n_0 + \alpha_1 n_0) k_1^2 \Phi_1^{*(2)} + (1 - n_0 + \alpha_2 n_0) k_2^2 \Phi_2^{*(2)} \right] \quad (22)$$

where wave numbers k_j and solid/fluid related parameters α_j are given as (8)–(20) in Huang and Song (1993). In (21), $\Phi_1^{*(2)}$ and $\Phi_2^{*(2)}$ are displacement potentials of the first kind and the second kind of longitudinal waves, respectively, and $\Phi_3^{*(2)}$ is the displacement potential of the third kind of transverse wave; that is

$$\mathbf{d}^* = \nabla\Phi_1^{*(2)} + \nabla\Phi_2^{*(2)} + \nabla^{\wedge}(\Phi_3^{*(2)}\mathbf{e}_z) \quad (23)$$

$$\mathbf{D}^* = \alpha_1\nabla\Phi_1^{*(2)} + \alpha_2\nabla\Phi_2^{*(2)} + \alpha_3\nabla^{\wedge}(\Phi_3^{*(2)}\mathbf{e}_z) \quad (24)$$

Note that the governing equations (1) and (21) and the pressure and effective stresses (2), (22), and (6), together with boundary conditions (14)–(17), (19), and (20), form the complete boundary-value problem of the present study.

Nondimensionalized Governing Equations and Boundary Conditions

Huang and Song (1993) defined the following parameters in their solution of a linear water wave propagating over a poroelastic bed:

$$m = (2G + \lambda)n_0/K \quad (25)$$

$$\varepsilon = n_0\rho_0\omega/b \quad (26)$$

$$\Lambda^2 = \frac{n_0\rho_0 + (1 - n_0)\rho_s}{2G + \lambda + (K/n_0)} \frac{\omega^2}{k_0^2} \quad (27)$$

$$\Pi^2 = \frac{i(m + 1)}{m\varepsilon} \frac{\rho_0}{K} \frac{\omega^2}{k_0^2} \quad (28)$$

$$\Psi^2 = \frac{n_0\rho_0 + (1 - n_0)\rho_s}{G} \frac{\omega^2}{k_0^2} \quad (29)$$

in which ε = penetrability parameter; ω = frequency; m = stiffness ratio of solid and fluid; Λ and Ψ are only functions of water wave speed and material (fluid and solid skeleton) properties, whereas Π is not only a function of the same variables for Λ and Ψ but also depends on the permeability of porous medium.

For low penetrability (i.e., $\|\varepsilon\| \ll 1$), (27)–(29) could be simplified to

$$\Lambda^2 \cong (k_1/k_0)^2 \quad (30)$$

$$\Pi^2 \cong (k_2/k_0)^2 \quad (31)$$

$$\Psi^2 \cong (k_3/k_0)^2 \quad (32)$$

Moreover, for soft solid skeleton, $\|k_2\| \gg \|k_0\|$ and such that $\|\Pi^2\| \gg 1$. Because $\|\Lambda^2\|$ is always smaller than $\|\Psi^2\|$, one can obtain

$$\|\Lambda^2\| < \|\Psi^2\| \ll 1 \ll \|\Pi^2\| \quad (33)$$

After the analysis of order of magnitude for each dependent variable, the dimensionless variables are selected as

$$\hat{x} = k_0x \quad (34)$$

$$\hat{y} = k_0y \quad (35a)$$

$$y' = \hat{y}/\varepsilon_2 \quad \text{for the second longitudinal wave only} \quad (35b)$$

$$\hat{t} = \sqrt{gk_0}t \quad (36)$$

$$\hat{\eta}^* = k_0\eta^* \quad (37)$$

$$\hat{\omega} = \omega/\sqrt{gk_0} \quad (38)$$

$$\hat{\Phi}^{*(1)} = \frac{k_0^2}{\sqrt{gk_0}} \Phi^{*(1)} \quad (39)$$

$$\hat{\Phi}_1^{*(2)} = e^{k_0h}k_0^2\Phi_1^{*(2)} \quad (40)$$

$$\Phi_2^{*(2)} = e^{k_0h}k_0^2 \frac{k_2^2}{k_1^2} \Phi_2^{*(2)} = \frac{e^{k_0h}k_0^2}{\varepsilon_2^2\Lambda^2} \Phi_2^{*(2)} \quad (41)$$

$$\Phi_3^{*(2)} = e^{k_0h}k_0^2\Phi_3^{*(2)} \quad (42)$$

$$\hat{\xi}^* = \frac{k_0e^{k_0h}}{\Psi^2} \xi^* \quad (43)$$

$$\hat{p}^{*(1)} = \frac{k_0}{\rho_0g} \hat{p}^{*(1)} \quad (44)$$

$$\hat{p}^{*(2)} = \frac{k_0}{\rho_0g} \hat{p}^{*(2)} \quad (45)$$

All of the symbols of variables on the left-hand side of (34)–(45) are dimensionless, but those on the right-hand side are dimensional. Note that because the vertical length scales need multiple scales, \hat{y} and y' are proposed for the boundary layer correction approach.

Applying the two-parameter perturbation expansion, velocity potential of channel flow and displacement potentials of the three kinds of wave for the whole domain can be written

$$\hat{\Phi}^{*(1)} = \varepsilon_1\hat{\Phi}_{10}^* + \varepsilon_1\varepsilon_2\hat{\Phi}_{11}^* + \varepsilon_1^2\hat{\Phi}_{20}^* + O(\varepsilon_1^2\varepsilon_2, \dots) \quad (46)$$

$$\hat{\Phi}_j^{*(2)} = \varepsilon_1\hat{\Phi}_{10}^{*[j]} + \varepsilon_1\varepsilon_2\hat{\Phi}_{11}^{*[j]} + \varepsilon_1^2\hat{\Phi}_{20}^{*[j]} + O(\varepsilon_1^2\varepsilon_2, \dots), \quad j = 1, 3 \quad (47)$$

Due to (33), the second kind of wave needs to be solved inside the boundary layer. The displacement potential of the second kind of wave inside the boundary layer is nondimensionalized specifically as (41) with a magnified length scale [see (35b)] and expanded as

$$\hat{\Phi}_2^{*(2)} = \varepsilon_1\hat{\Phi}_{10}^{*[2]} + \varepsilon_1\varepsilon_2\hat{\Phi}_{11}^{*[2]} + \varepsilon_1^2\hat{\Phi}_{20}^{*[2]} + O(\varepsilon_1^2\varepsilon_2, \dots) \quad (48)$$

if $\|\varepsilon_2\|$ and $\|\varepsilon_1\|$ are smaller than unity. Also, the water wave at the free surface becomes

$$\hat{\eta}^* = \varepsilon_1\hat{\eta}_{10}^* + \varepsilon_1\varepsilon_2\hat{\eta}_{11}^* + \varepsilon_1^2\hat{\eta}_{20}^* + O(\varepsilon_1^2\varepsilon_2, \dots) \quad (49)$$

and the wave of the channel-bed interface becomes

$$\hat{\xi}^* = \varepsilon_1\hat{\xi}_{10}^* + \varepsilon_1\varepsilon_2\hat{\xi}_{11}^* + \varepsilon_1^2\hat{\xi}_{20}^* + O(\varepsilon_1^2\varepsilon_2, \dots) \quad (50)$$

For a periodic motion with frequency ω , the aforementioned variables $[\]^*(\mathbf{R}, t)$ can be written as $[\](\mathbf{R})e^{-i\omega t}$, where \mathbf{R} is a position vector. Let the given incoming wave amplitude before being disturbed by the porous bed be a , and if the wave number of this incoming wave is found to be k_0 (it will be found as complex), the Stokes expansion of two parameters based on ε_1 and ε_2 will be carried out only to the first three terms of the present nonlinear water wave problem to avoid the occurrence of secular terms. Thus, after the Taylor series expansions at the free surface and at the channel-bed interface, respectively, the boundary-value problem of each order, without the time factor, is obtained in the following.

Boundary-Value Problem for the First and Third Kinds of Wave

$O(\varepsilon_1)$

Governing equations

- Region 1: $-\infty < \hat{x} < \infty, 0 < \hat{y} < k_0h$

$$\hat{\nabla}^2\hat{\Phi}_{10} = 0 \quad (51)$$

- Region 2: $-\infty < \hat{x} < \infty, -\infty < \hat{y} < 0$

$$\hat{\nabla}^2\hat{\Phi}_{10}^{[3]} + \Psi^2\hat{\Phi}_{10}^{[3]} = 0 \quad (53)$$

Boundary conditions

- At the free surface: $\hat{y} = k_0 h$, $-\infty < \hat{x} < \infty$
(a) Kinematic free surface boundary condition

$$\hat{\phi}_{10,\hat{y}} = -i\hat{\omega}\hat{\eta}_{10} \quad (54)$$

- (b) Dynamic free surface boundary condition

$$-i\hat{\omega}\hat{\phi}_{10} + \hat{\eta}_{10} = 0 \quad (55)$$

- At the porous bed interface: $\hat{y} = 0$, $-\infty < \hat{x} < \infty$
(a) Continuity of pressure

$$-i\hat{\omega}\hat{\phi}_{10} + \frac{k_0 K \Lambda^2}{e^{k_0 h} n_0 \rho_0 g} q_1 \hat{\phi}_{10}^{[1]} = 0 \quad (56)$$

- (b) Continuity of flux

$$i\hat{\omega} q_1 \hat{\phi}_{10,\hat{y}}^{[1]} - i\hat{\omega} q_3 \hat{\phi}_{10,\hat{x}}^{[3]} + e^{k_0 h} \hat{\phi}_{10,\hat{y}} = 0 \quad (57)$$

- (c) Continuity of effective stress (only $\tau_{\hat{x}\hat{y}} \doteq 0$)

$$2\hat{\phi}_{10,\hat{x}\hat{y}}^{[1]} + \hat{\phi}_{10,\hat{y}\hat{y}}^{[3]} - \hat{\phi}_{10,\hat{x}\hat{x}}^{[3]} = 0 \quad (58)$$

- At the deep far field: $\hat{y} \rightarrow -\infty$, $-\infty < \hat{x} < \infty$

$$\hat{\phi}_{10}^{[j]} \rightarrow 0, \quad j = 1, 3 \quad (59)$$

where

$$q_j = 1 - n_0 + \alpha_j n_0, \quad j = 1, 3 \quad (60)$$

Note that only one component of the above boundary condition of continuity of effective stress is needed (i.e., $\tau_{\hat{x}\hat{y}} \doteq 0$), otherwise it will become overdetermined. (Another condition, $\tau_{\hat{y}\hat{y}} \doteq 0$, includes the effect of the second kind of wave and will be adopted by the boundary layer correction for the second kind of wave later.)

$O(\varepsilon_1 \varepsilon_2)$

The presentations are identical with those in $O(\varepsilon_1)$ except subscripts are changed from 10 to 11.

$O(\varepsilon_1^2)$

Governing equations

- Region 1: $-\infty < \hat{x} < \infty$, $0 < \hat{y} < k_0 h$

$$\hat{\nabla}^2 \hat{\phi}_{20} = 0 \quad (61)$$

- Region 2: $-\infty < \hat{x} < \infty$, $-\infty < \hat{y} < 0$

$$\hat{\nabla}^2 \hat{\phi}_{20}^{[1]} + \frac{\tilde{k}_1^2}{k_1^2} \Lambda^2 \hat{\phi}_{20}^{[1]} = 0 \quad (62)$$

$$\hat{\nabla}^2 \hat{\phi}_{20}^{[3]} + \frac{\tilde{k}_3^2}{k_3^2} \Psi^2 \hat{\phi}_{20}^{[3]} = 0 \quad (63)$$

Boundary conditions

- At the free surface: $\hat{y} = k_0 h$, $-\infty < \hat{x} < \infty$
(a) Kinematic free surface boundary condition

$$\hat{\phi}_{20,\hat{y}} + 2i\hat{\omega}\hat{\eta}_{20} = \hat{\eta}_{10,\hat{x}} \hat{\phi}_{10,\hat{x}} - \hat{\eta}_{10} \hat{\phi}_{10,\hat{y}\hat{y}} \quad (64)$$

- (b) Dynamic free surface boundary condition

$$-2i\hat{\omega}\hat{\phi}_{20} + \hat{\eta}_{20} = i\hat{\omega}\hat{\eta}_{10}\hat{\phi}_{10,\hat{y}} - \frac{1}{2}(\hat{\phi}_{10,\hat{x}}^2 + \hat{\phi}_{10,\hat{y}}^2) \quad (65)$$

- At the porous bed interface: $\hat{y} = 0$, $-\infty < \hat{x} < \infty$
(a) Continuity of pressure

$$2i\hat{\omega}\hat{\phi}_{20} - \frac{k_0 K \Lambda^2}{e^{k_0 h} n_0 \rho_0 g} \tilde{q}_1 \hat{\phi}_{20}^{[1]} = \frac{1}{2}(\hat{\phi}_{10,\hat{x}}^2 + \hat{\phi}_{10,\hat{y}}^2) - \frac{i\hat{\omega}\Psi^2 \tilde{\xi}_{10}^2}{e^{k_0 h}} \hat{\phi}_{10,\hat{y}} + \frac{\Psi^2 k_0 K \Lambda^2 \tilde{\xi}_{10}^2}{e^{2k_0 h} n_0 \rho_0 g} q_1 \hat{\phi}_{10,\hat{y}}^{[1]} \quad (66)$$

- (b) Continuity of flux

$$e^{k_0 h} \hat{\phi}_{20,\hat{y}} + 2i\hat{\omega}(\tilde{q}_1 \hat{\phi}_{20,\hat{y}}^{[1]} - \tilde{q}_3 \hat{\phi}_{20,\hat{x}}^{[3]}) = \Psi^2 \tilde{\xi}_{10,\hat{x}} \hat{\phi}_{10,\hat{x}} - \Psi^2 \tilde{\xi}_{10} \hat{\phi}_{10,\hat{y}\hat{y}} + i\hat{\omega}\Psi^2 e^{-k_0 h} \tilde{\xi}_{10,\hat{x}} (q_1 \hat{\phi}_{10,\hat{x}}^{[1]} + q_3 \hat{\phi}_{10,\hat{y}}^{[3]}) - i\hat{\omega}\Psi^2 e^{-k_0 h} \tilde{\xi}_{10} (q_1 \hat{\phi}_{10,\hat{y}\hat{y}}^{[1]} - q_3 \hat{\phi}_{10,\hat{x}\hat{x}}^{[3]}) \quad (67)$$

- (c) Continuity of effective stress (only $\tau_{\hat{x}\hat{y}} \doteq 0$)

$$G(2\hat{\phi}_{20,\hat{x}\hat{y}}^{[1]} + \hat{\phi}_{20,\hat{y}\hat{y}}^{[3]} - \hat{\phi}_{20,\hat{x}\hat{x}}^{[3]}) = \Psi^2 e^{-k_0 h} \tilde{\xi}_{10,\hat{x}} [2G(\hat{\phi}_{10,\hat{x}\hat{x}}^{[1]} + \hat{\phi}_{10,\hat{y}\hat{y}}^{[3]}) - \lambda \Lambda^2 \hat{\phi}_{10}^{[1]} - G\Psi^2 e^{-k_0 h} \tilde{\xi}_{10} (2\hat{\phi}_{10,\hat{x}\hat{y}}^{[1]} + \hat{\phi}_{10,\hat{y}\hat{y}}^{[3]} - \hat{\phi}_{10,\hat{x}\hat{x}}^{[3]})] \quad (68)$$

- At the deep far field: $\hat{y} \rightarrow -\infty$, $-\infty < \hat{x} < \infty$

$$\hat{\phi}_{20}^{[j]} \rightarrow 0, \quad j = 1, 3 \quad (69)$$

where \tilde{k}_j and $\tilde{\alpha}_j$ ($j = 1, 3$) in nonlinear order ε_1^2 are given as eqs. (8)–(20) in Huang and Song (1993), and

$$\tilde{q}_j = 1 - n_0 + \tilde{\alpha}_j n_0, \quad j = 1, 3 \quad (70)$$

Boundary Layer Correction for Second Kind of Wave

The second kind of wave disappears outside the boundary layer, but it does exist inside the boundary layer; thus the complete solution needs to be corrected by further considering the second kind of wave inside the porous material bed. Because a thin boundary layer exists within the porous bed near the water/porous bed interface, multiple scales are necessary to solve the nonlinear boundary-value problem for the second longitudinal wave. One therefore lets $y' = \hat{y}/\varepsilon_2$ to change the scale from \hat{y} to the magnified scale y' in (35b). The difficulty (i.e., the error due to the first partial derivative based on y of the displacement potential of the second longitudinal wave) that Chen et al. (1997) encountered is now overcome by proposing this double length scale in the vertical direction. After the coordinate transformation of (35b), the boundary-value problem of the displacement potential of the second longitudinal wave inside the boundary layer is as follows.

$O(\varepsilon_1)$

Governing equations

$$\hat{\phi}_{10,y'y'}^{[2]} + \hat{\phi}_{10}^{[2]} = 0 \quad (71)$$

Boundary Conditions: $y' = 0$, $-\infty < \hat{x} < \infty$

- (a) Continuity of pressure

$$\hat{\phi}_{10}^{[2]} = 0 \quad (72)$$

- (b) Continuity of effective stress (only $\tau_{\hat{y}\hat{y}} \doteq 0$)

$$2G\Lambda^2 \hat{\phi}_{10,y'y'}^{[2]} - \lambda \Lambda^2 \hat{\phi}_{10}^{[2]} = \lambda \Lambda^2 \hat{\phi}_{10}^{[1]} - 2G(\hat{\phi}_{10,\hat{y}\hat{y}}^{[1]} - \hat{\phi}_{10,\hat{x}\hat{x}}^{[3]}) \quad (73)$$

Note that the boundary condition of continuity of flux is equivalent to (72), and the boundary condition of $\tau_{\hat{y}\hat{y}} \doteq 0$ just satisfies $\hat{\phi}_{10,\hat{x}\hat{y}}^{[2]} = -\hat{\phi}_{10,y'\hat{x}}^{[2]}$ (i.e., $\hat{\phi}_{10,\hat{x}\hat{y}}^{[2]} = 0$). However, $\hat{\phi}_{10,\hat{x}\hat{y}}^{[2]} = 0$ at $y' = 0$ is automatically satisfied by referring to (72). Thus only

one component of the effective stress boundary condition $\tau_{yy} \equiv 0$ is needed to solve $\hat{\phi}_{10}^{[2]}$.

$O(\epsilon_1 \epsilon_2)$

The presentations are identical with those in $O(\epsilon_1)$ except subscripts are changed from 10 to 11.

$O(\epsilon_1^2)$

Governing equations

$$\hat{\phi}_{20,y'y'}^{[2]} + \hat{\phi}_{20}^{[2]} = 0 \quad (74)$$

Boundary conditions: $y' = 0, -\infty < \hat{x} < \infty$

(a) Continuity of pressure

$$\begin{aligned} \frac{k_0 K \Lambda^2}{e^{k_0 h} n_0 \rho_0 g} \tilde{q}_2 \hat{\phi}_{20}^{[2]} &= 2i\hat{\omega} \hat{\phi}_{20} - \frac{1}{2} (\hat{\phi}_{10,x}^{[2]} + \hat{\phi}_{10,y}^{[2]}) \\ &- \frac{k_0 K \Lambda^2}{e^{k_0 h} n_0 \rho_0 g} \tilde{q}_1 \hat{\phi}_{20}^{[2]} - \Psi^2 \hat{\xi}_{10} e^{-k_0 h} \\ &\cdot \left[-i\hat{\omega} \hat{\phi}_{10,y} + \frac{k_0 K \Lambda^2}{e^{k_0 h} n_0 \rho_0 g} (q_1 \hat{\phi}_{10,y}^{[1]} + q_2 \hat{\phi}_{11,y}^{[2]}) \right] \end{aligned} \quad (75)$$

(b) Continuity of effective stress (only $\tau_{yy} \equiv 0$)

$$\begin{aligned} 2G\Lambda^2 \hat{\phi}_{20,y'y'}^{[2]} - \lambda\Lambda^2 \hat{\phi}_{20}^{[2]} &= -2G(\hat{\phi}_{20,y'y}^{[1]} - \hat{\phi}_{20,y'y}^{[3]}) + \lambda\Lambda^2 \hat{\phi}_{20}^{[1]} \\ &- \Psi^2 \hat{\xi}_{10} e^{-k_0 h} [2G(\hat{\phi}_{20,y'y}^{[1]} + \Lambda^2 \hat{\phi}_{11,y'y'}^{[2]} - \hat{\phi}_{10,y'y}^{[3]}) \\ &- \lambda\Lambda^2 (\hat{\phi}_{10,y}^{[1]} + \hat{\phi}_{11,y}^{[2]})] \end{aligned} \quad (76)$$

SOLUTION

After omitting the time factor $e^{-i\omega t}$, the given incoming water wave profile with magnitude a is

$$\eta_{10}(x) = ae^{ik_0 x}, \quad 0 < x < \infty \quad (77)$$

With the input of the incoming water wave, each order of the aforementioned boundary-value problem can be solved in sequence as shown in Fig. 2. Fig. 2 indicates that one can first find the solution of order ϵ_1 outside the boundary layer and can subsequently match it with the inner expansion to complete the solution of order ϵ_1 . Then, the solution of order ϵ_1 is provided to solve the higher orders in sequence as indicated by the solid arrows as shown in Fig. 2. Note that the problem of order $\epsilon_1 \epsilon_2$ is solved simultaneously to find the unique solution. Thus the dimensional solutions of the first kind of longitudinal wave and the third kind of transverse wave throughout the entire domain are obtained as follows.

$O(\epsilon_1)$

$$\phi_{10} = -\frac{i}{k_0} \left[\frac{g}{\omega} \cosh k_0(h-y) - \frac{\omega}{k_0} \sinh k_0(h-y) \right] e^{ik_0 x} \quad (78)$$

$$\hat{\phi}_{10}^{[1]} = \frac{i}{k_0^2 (iq_1 K_1 - q_3 L_1)} \left[\cosh(k_0 h) - \frac{gk_0}{\omega^2} \sinh(k_0 h) \right] e^{K_1 k_0 y + ik_0 x} \quad (79)$$

$$\hat{\phi}_{10}^{[3]} = \frac{a_3}{e^{k_0 h} k_0^2} e^{K_3 k_0 y + ik_0 x} \quad (80)$$

with the dispersion relation of complex wave number k_0 as

$$T_1 - 1 = \left(\frac{T_1 \omega^2}{gk_0} - \frac{gk_0}{\omega^2} \right) \tanh(k_0 h) \quad (81)$$

where

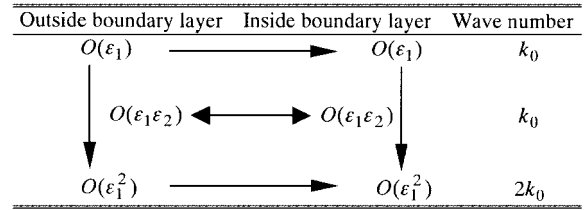


FIG. 2. Relation among Solutions of Each Order

$$T_1 = \frac{n_0 \rho_0 g (q_1 K_1 + iq_3 L_1)}{k_0 K \Lambda^2 q_1} \quad (82)$$

$O(\epsilon_1 \epsilon_2)$

$$\phi_{11} = \frac{\sqrt{gk_0}}{k_0^2} E_5 \left[\cosh(k_0 y) - \frac{\omega^2}{gk_0} \sinh(k_0 y) \right] e^{ik_0 x} \quad (83)$$

$$\phi_{11}^{[1]} = \frac{c_1}{e^{k_0 h} k_0^2} e^{K_1 k_0 y + ik_0 x} \quad (84)$$

$$\phi_{11}^{[3]} = \frac{c_3}{e^{k_0 h} k_0^2} e^{K_3 k_0 y + ik_0 x} \quad (85)$$

$$\eta_{11} = \frac{i}{k_0} \frac{\omega}{\sqrt{gk_0}} E_5 e^{ik_0 x} \quad (86)$$

$O(\epsilon_1^2)$

$$\phi_{20} = \frac{\sqrt{gk_0}}{k_0^2} [E_3 \cosh 2k_0(h-y) + E_4 \sinh 2k_0(h-y)] e^{2ik_0 x} \quad (87)$$

$$\phi_{20}^{[1]} = \frac{b_1}{e^{k_0 h} k_0^2} e^{M_1 k_0 y + 2ik_0 x} \quad (88)$$

$$\phi_{20}^{[3]} = \frac{b_3}{e^{k_0 h} k_0^2} e^{M_3 k_0 y + 2ik_0 x} \quad (89)$$

$$\eta_{20} = \left(\frac{g}{\omega^2} - \frac{iE_4}{\omega k_0} \sqrt{gk_0} \right) e^{2ik_0 x} \quad (90)$$

The dimensional solutions of the second kind of longitudinal wave obtained by the boundary layer correction approach are as follows.

$O(\epsilon_1)$

$$\phi_{10}^{[2]} = \frac{a_2}{k_0^2 e^{k_0 h} k_0^2} \exp \left[ik_0 \left(x - \frac{y}{\epsilon_2} \right) \right] \quad (91)$$

with

$$a_2 = (C_0 t_1 - q_1 a_1) / q_2 \quad (92)$$

$O(\epsilon_1 \epsilon_2)$

$$\phi_{11}^{[2]} = \frac{c_2}{k_0^2 e^{k_0 h} k_0^2} \exp \left[ik_0 \left(x - \frac{y}{\epsilon_2} \right) \right] \quad (93)$$

with

$$c_2 = \left(i \frac{\omega}{\sqrt{gk_0}} C_0 t_1 E_5 - q_1 c_1 \right) / q_2 \quad (94)$$

$O(\epsilon_1^2)$

$$\phi_{20}^{[2]} = \frac{b_2}{k_0^2 e^{k_0 h} k_0^2} \exp \left[ik_0 \left(2x - \frac{y}{\epsilon_2} \right) \right] \quad (95)$$

with

$$b_2 = e^{-k_0 h} \tilde{q}_2^{-1} \left\{ 2i \frac{\omega}{\sqrt{gk_0}} r_4 C_0 e^{2k_0 h} - \frac{C_0 g k_0}{2\omega^2} (t_1^2 - r_1^2) + (K_1 a_1 - ia_3)[C_0 r_1 - (q_1 a_1 K_1 - iq_2 c_2)] - e^{k_0 h} \tilde{q}_1 b_1 \right\} \quad (96)$$

After solving the displacement potentials, all of the other variables can be obtained. The wave of the porous bed interface from (18) gives

$$\xi(x) = a(K_1 a_1 - ia_3) e^{ik_0 x - k_0 h} + \varepsilon_2 a e^{ik_0 x - k_0 h} (c_1 K_1 - ia_2 \Lambda^2 - ic_3) + \varepsilon_1 \frac{a}{2} e^{2ik_0 x - 2k_0 h} \{ (K_1 a_1 - ia_3)[(K_1^2 + 1)a_1 - a_2 \Lambda^2 - 2iK_3 a_3] + 2b_1 M_1 e^{k_0 h} - 4ib_3 e^{k_0 h} \} \quad (97)$$

RESULTS AND COMMENTS

Because the conventional higher-order Stokes expansion based on one parameter, $\varepsilon_1 = k_0 a$, breaks down for the soft material (Chen et al. 1997), this study proposes a two-parameter expansion with ε_1 and $\varepsilon_2 = k_0/k_2$ to estimate the second kind of longitudinal wave for the soft material. In addition, one finds that when the seabed is soft, the solution of Huang and Song (1993) is constrained to $\|\varepsilon_2\| > \|\varepsilon_1\|$ with $n \geq 1$ (i.e., $\|\varepsilon_2^n \varepsilon_1\| > \|\varepsilon_1^2\|$), but the present method is only constrained to $\|\varepsilon_1\| > \|\varepsilon_2\|$ and $\|\varepsilon_2\| > \|\varepsilon_2^n\|$ (referring to Fig. 3 and comparing layer 2 with layer 3). In other words, the present solutions are valid when $\|\varepsilon_1\|^{1/2} > \|\varepsilon_2\| > \|\varepsilon_1\|^2$ before shear failure, which

indicates that the present solution is more accurate than the solution of Huang and Song (1993) in the same order of ε_2 for soft material. This is indicated clearly in Fig. 3, the schematic diagram of the two-parameter expansion. Because the second longitudinal wave decays very quickly in the y-direction near the homogeneous water/porous bed interface, Chen et al. (1997) failed to estimate accurately the first order of partial derivative with respect to y for the displacement potential of the second longitudinal wave inside the boundary layer

TABLE 1. Material Properties of Very Soft Soil

Item (1)	Notation (2)	Value (3)	Unit (4)
(a) Water			
Density	ρ_0	1,000	kg/m ³
Bulk modulus	K	2.3×10^9	N/m ²
Dynamic viscosity	μ	0.001	Ns/m ²
Depth	h	0.35	m
Wave height	a	0.0357	m
Period	T	1.55	s
(b) Skeleton			
Density	ρ_s	2,650	kg/m ³
Lame's constant	G	5.0×10^4	N/m ²
Lame's constant	λ	1.0×10^5	N/m ²
Specific permeability	k_p	1.0×10^{-11}	m ²
Porosity	n_0	0.4	—

TABLE 2. Variables of Very Soft Soil

Notation (1)	Value (2)
k_0	0.245834E+01+i0.481295E-04
k_1	0.238469E-02+i0.237891E-07
k_2	0.318363E+02+i0.318331E+02
k_3	0.808703E+00+i0.823670E-05
\tilde{k}_1	0.476939E-02+i0.952280E-07
\tilde{k}_2	0.450257E+02+i0.450166E+02
\tilde{k}_3	0.161741E+01+i0.329468E-04
a_1	0.127676E+01-i0.674232E-03
a_2	0.386868E-01-i0.221286E+02
a_3	-0.714250E-03-i0.134980E+01
b_1	0.171970E+01+i0.707839E-03
b_2	-0.118767E-03-i0.341805E+00
b_3	0.749158E-03-i0.181747E+01
c_1	0.208305E-04-i0.189383E-04
c_2	0.110643E+02+i0.284844E-02
c_3	-0.200597E-04-i0.836553E-08

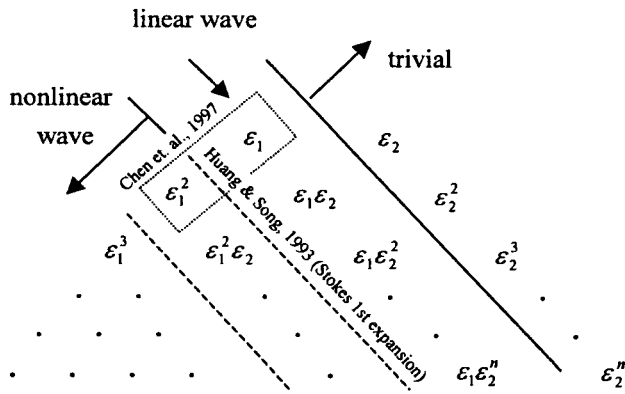


FIG. 3. Schematic Diagram of Two-Parameter Expansion

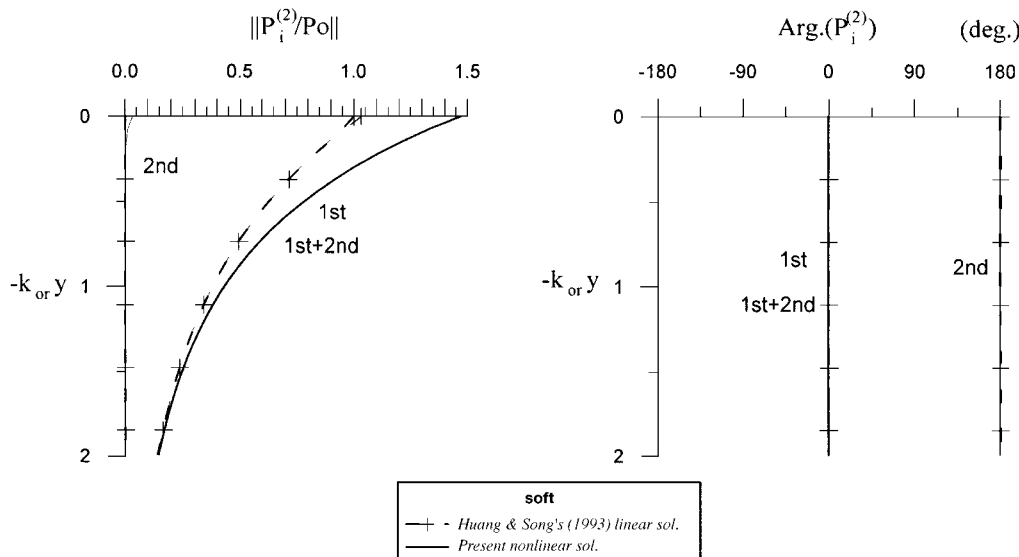


FIG. 4. Variation of Pore Pressure for Very Soft Material ($\|\varepsilon_1\| = 0.0878$, $\|\varepsilon_2\| = 0.0546$)

$\hat{\Phi}_2^{*(2)}$. Therefore, one more length scale is needed to estimate more accurately the first order of partial derivative for $\hat{\Phi}_2^{*(2)}$ in the vertical direction by the transformation of y to y' [see (35b)]. Because $\Delta y'$ is much shorter than Δy , the length scale adopted by Chen et al. (1997) will overestimate the displacement potential of the second longitudinal wave. That is why the present method adopts two length scales to reformulate the boundary-value problem inside the boundary layer.

To confirm the validity of the proposed solution, some soft material and wave conditions are selected to compute pore

pressure and effective stresses inside the porous bed as well as the profiles of the channel bed and free surface. The results agree very well to the solutions of Huang and Song (1993). (Because the results are so close to those of Huang and Song's solutions, the detailed figures will not be shown here to save space.) Because there is good agreement between the present solutions and those of Huang and Song (1993) for the soft porous bed as discussed, one may simply use the present solution to show the validity of Biot's theory of poroelasticity (Biot 1962) in simulating a soft porous bed.

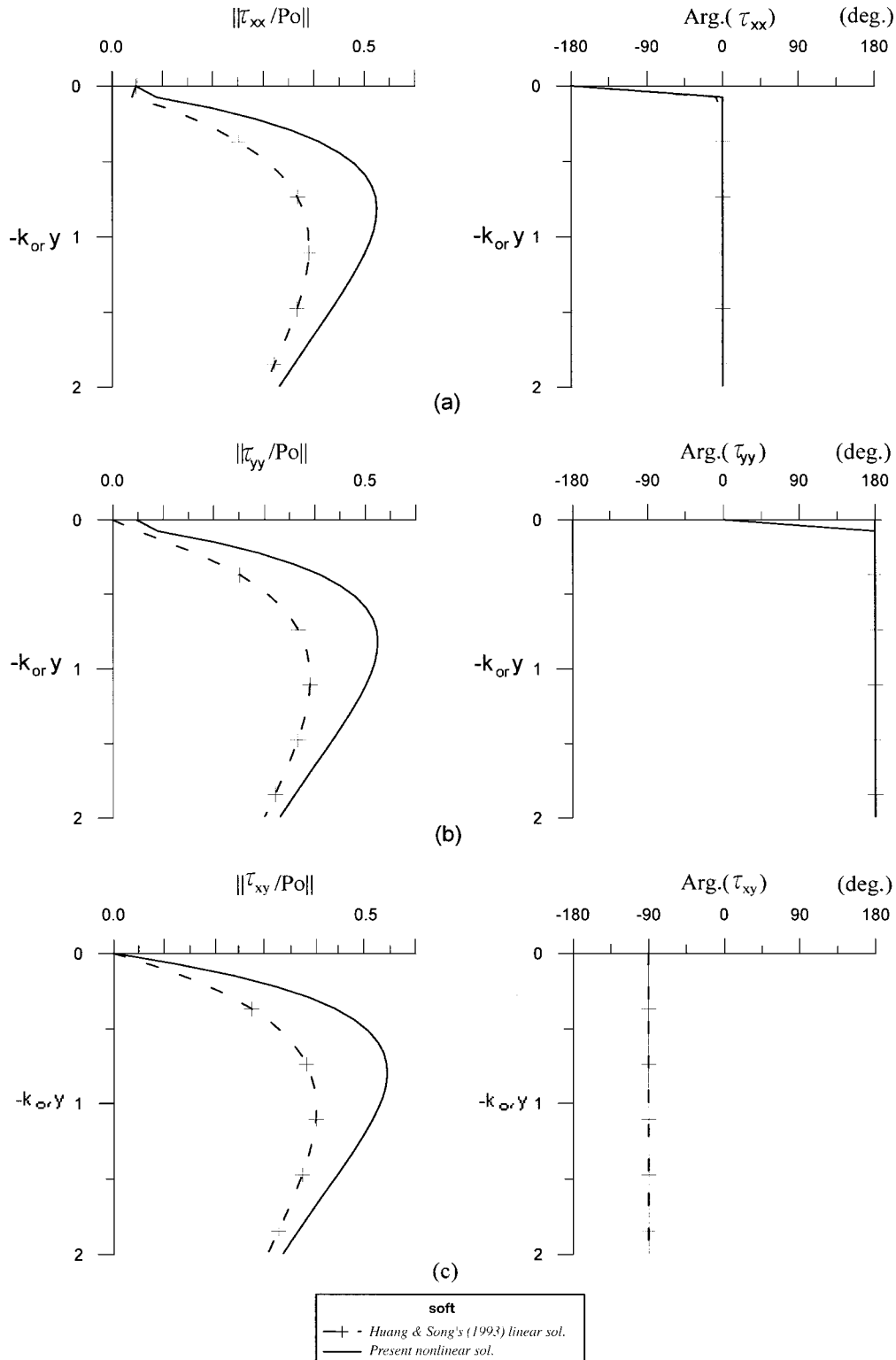


FIG. 5. Variation of Effective Stress for Very Soft Material: (a) xx -component; (b) yy -component; (c) xy -component ($\|\varepsilon_1\| = 0.0878$, $\|\varepsilon_2\| = 0.0546$)

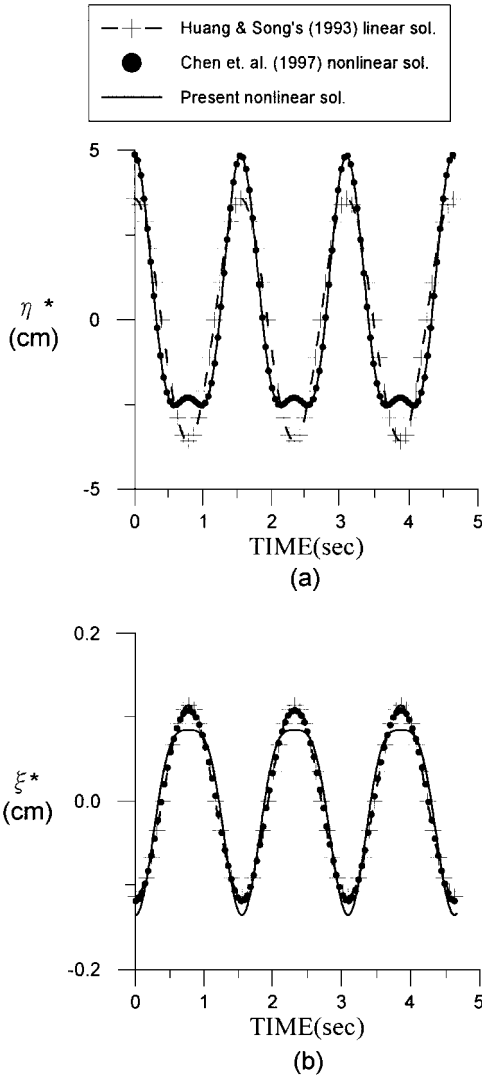


FIG. 6. (a) Free Surface Profiles; (b) Channel-Bed Interface Profiles of Wave Period 1.55 s with Respect to Time for Very Soft Material ($\|\varepsilon_1\| = 0.0878$, $\|\varepsilon_2\| = 0.0546$)

Table 1 shows the material properties of very soft soil, and Table 2 provides the simulation result. Comparing the present Table 2 with Table 2 of Chen et al. (1997), one can find that the unreasonable result—the coefficient of a higher-order term is much larger than that of the leading term (i.e., $|b_2| \gg |a_2|$)—does not exist in the present result.

The very soft soil case is illustrated to show the differences between the linear solution of Huang and Song (1993) and the present nonlinear solution, and the simulation results are plotted in Figs. 4–6. In Fig. 4, P_0 is the perturbed pressure on bed at $y = 0$. In fact, the analytic solution proposed by Huang and Song (1993) failed in this case due to their constrain, $\|\varepsilon_2'\| > \|\varepsilon_1\|$, but the present method is still applicable. Note that the Ursell number, $U_r = \text{Re}(k_0 a) / [\text{Re}(k_0 h)]^3$, is 0.138. Moreover, referring to Fig. 6(a) the water profile is the same as that of Chen et al. (1997) while referring to Fig. 6(b), the soil profile is significantly different than that of Chen et al. (1997). This is because the solution of homogeneous water of Chen et al. (1997) is correct, but the solution of the second longitudinal wave inside the porous medium is wrong. Furthermore, referring to Figs. 7(b–d) of Chen et al. (1997) the experimental data near the peaks and troughs are lower than their nonlinear solution, and the present solution in Fig. 6(b) also has the same tendency as the experimental data [i.e., the present nonlinear solution is lower than that of Chen et al. (1997) in the peaks

and the troughs]. Thus combining Figs. 4–6, it can be concluded that in simulating a soft porous bed it is proper to adopt a two-parameter expansion instead of the conventional one-parameter expansion.

CONCLUSIONS

When the bed material is soft, the higher-order Stokes expansion of a water wave based on ε_1 is invalid [i.e., the one-parameter perturbation failed (Chen et al. 1997)]. This is because a boundary layer exists inside the porous bed and near the homogeneous water/porous bed interface from the second longitudinal wave. Therefore, considering that the second length scale based on the second longitudinal wave number for the boundary-value problem is necessary, a two-parameter perturbation expansion based on ε_1 and ε_2 is proposed. Because the second kind of longitudinal wave vanishes outside the boundary layer, but exists inside the boundary layer, the complete solution of the displacement potential needs to be corrected inside the boundary layer. Hence the complete boundary-value problem is thus solved by the boundary layer correction approach in the present study. Referring to Chen et al. (1997), a nonlinear water wave is very likely to happen even when $\|\varepsilon_1\|$ and the Ursell parameter are very small, and this is proved again in the present study.

Moreover, due to the effect of the second longitudinal wave inside the soft porous bed, it is very difficult to compute the pore pressure, effective stresses, and bed form accurately near the interface by the numerical model. The numerical model will encounter a problem of convergence even by finding the meshes near the interface. Thus the concept of the present study is very helpful in formulating a simplified boundary-value problem in numerical computation for a soft poroelastic bed with irregular geometry.

APPENDIX I. COEFFICIENTS OF SOLUTIONS OUTSIDE BOUNDARY LAYER

$O(\varepsilon_1)$

$$a_3 = -L_1 a_1 \quad (98)$$

$$a_1 = \frac{ie^{k_0 h}}{iq_1 K_1 - q_3 L_1} \left[\cosh(k_0 h) - \frac{gk_0}{\omega^2} \sinh(k_0 h) \right] \quad (99)$$

$$L_1 = \frac{2iK_1}{1 + K_3^2} \quad (100)$$

$$K_1^2 = 1 - \frac{k_1^2}{k_0^2} \quad (101)$$

$$K_3^2 = 1 - \frac{k_3^2}{k_0^2} \quad (102)$$

$O(\varepsilon_1 \varepsilon_2)$

$$E_3 = \frac{1}{t_2} \{ 2(iK_1 q_1 r_3 + r_2 q_3) - q_2 [2iK_1 r_3 + r_2 (K_3^2 + 1)] \} \cdot \frac{\Lambda^2}{q_2} (C_0 t_1 - q_1 a_1) \frac{\omega}{\sqrt{gk_0}} \quad (103)$$

$$c_1 = \frac{1}{t_2} \left\{ i[2q_3 - q_2 (K_3^2 + 1)] C_0 t_1 \frac{\omega^2}{gk_0} - 2r_1 r_3 \right\} \frac{\Lambda^2}{q_2} (C_0 t_1 - q_1 a_1) \quad (104)$$

$$c_3 = \frac{-2}{t_2} \left[C_0 K_1 t_1 \frac{\omega^2}{gk_0} (q_2 - q_1) + r_1 r_2 \right] \frac{\Lambda^2}{q_2} (C_0 t_1 - q_1 a_1) \quad (105)$$

$$t_2 = [2K_1 q_3 - K_1 q_3 (K_3^2 + 1)] C_0 t_1 \frac{\omega^2}{gk_0} + r_1 [2iK_1 r_3 + r_2 (K_3^2 + 1)] \quad (106)$$

$$t_1 = e^{k_0 h} \left[\cosh(k_0 h) - \frac{\omega^2}{gk_0} \sinh(k_0 h) \right] \quad (107)$$

$$r_1 = e^{k_0 h} \left[\frac{\omega^2}{gk_0} \cosh(k_0 h) - \sinh(k_0 h) \right] \quad (108)$$

$$r_2 = \frac{2GK_1^2 - \lambda\Lambda^2}{(2G + \lambda)\Lambda^2} q_2 + q_1 \quad (109)$$

$$r_3 = \frac{2iGK_3}{(2G + \lambda)\Lambda^2} q_2 \quad (110)$$

$$C_0 = \frac{n_0 \rho_0 g}{k_0 K \Lambda^2} \quad (111)$$

$$q_2 = 1 - n_0 + \alpha_2 n_0 \quad (112)$$

$O(\varepsilon_1^2)$

$$E_3 = \frac{f_{13} f_{22} f_{32} + f_{12} f_{23} f_{33} - f_{12} f_{24} f_{32} - f_{13} f_{23} f_{31}}{f_{11} f_{22} f_{32} - f_{12} f_{21} f_{32} - f_{11} f_{23} f_{31}} \quad (113)$$

$$E_4 = - \left[2 \frac{\omega^2}{gk_0} E_3 + \frac{3}{2} i \left(\frac{\sqrt{gk_0}}{\omega} - \frac{\omega^3}{gk_0 \sqrt{gk_0}} \right) \right] \quad (114)$$

$$b_1 = \frac{1}{f_{12}} (f_{13} - f_{11} E_3) \quad (115)$$

$$b_3 = \frac{1}{f_{32}} (f_{33} - f_{31} b_1) \quad (116)$$

$$f_{11} = 2i \frac{\omega}{\sqrt{gk_0}} \left[\cosh(2k_0 h) - 2 \frac{\omega^2}{gk_0} \sinh(2k_0 h) \right] \quad (117)$$

$$f_{12} = - \frac{k_0 K \Lambda^2 \tilde{q}_1}{e^{k_0 h} n_0 \rho_0 g} \quad (118)$$

$$f_{13} = \frac{1}{2} \left(\frac{gk_0}{\omega^2} - \frac{\omega^2}{gk_0} \right) - \frac{\omega}{e^{k_0 h} \sqrt{gk_0}} (K_1 a_1 - ia_3) \left[\frac{\omega}{\sqrt{gk_0}} \cosh(k_0 h) - \frac{\sqrt{gk_0}}{\omega} \sinh(k_0 h) \right] + \frac{k_0 K \Lambda^2}{e^{2k_0 h} n_0 \rho_0 g} (K_1 a_1 - ia_3) q_1 a_1 K_1 - 3 \left(1 - \frac{\omega^4}{g^2 k_0^2} \right) \sinh(2k_0 h) \quad (119)$$

$$f_{21} = 2 \frac{\omega^2}{gk_0} \cosh(2k_0 h) - \sinh(2k_0 h) \quad (120)$$

$$f_{22} = \frac{i}{e^{k_0 h}} \frac{\omega}{\sqrt{gh_0}} \tilde{q}_1 M_1 \quad (121)$$

$$f_{23} = \frac{2}{e^{k_0 h}} \frac{\omega}{\sqrt{gk_0}} \tilde{q}_3 \quad (122)$$

$$f_{24} = \frac{i}{e^{k_0 h}} (K_1 a_1 - ia_3) \left[\frac{\sqrt{gk_0}}{\omega} \cosh(k_0 h) - \frac{\omega}{\sqrt{gk_0}} \sinh(k_0 h) \right] - \frac{\omega}{2e^{2k_0 h} \sqrt{gk_0}} (K_1 a_1 - ia_3) [ia_1 q_1 (1 + K_1^2) + 2a_3 q_3 K_3] - \frac{3}{2} i \left(\frac{\sqrt{gk_0}}{\omega} - \frac{\omega^3}{gk_0 \sqrt{gk_0}} \right) \cosh(2k_0 h) \quad (123)$$

$$f_{31} = 4GM_1 \quad (124)$$

$$f_{32} = -iG(M_3^2 + 4) \quad (125)$$

$$f_{33} = e^{-k_0 h} (K_1 a_1 - ia_3) [2iGa_3 K_3 - (2G + \lambda\Lambda^2) a_1] + e^{-k_0 h} Gi (K_1 a_1 - ia_3) (2ia_1 K_1^2 + a_3 K_3^2 + a_3 K_3) \quad (126)$$

$$M_1^2 = 4 - \frac{\tilde{k}_1^2}{k_0^2} \quad (127)$$

$$M_3^2 = 4 - \frac{\tilde{k}_3^2}{k_0^2} \quad (128)$$

ACKNOWLEDGMENT

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APPENDIX II. REFERENCES

- Biot, M. A. (1962). "Mechanics of deformation and acoustic propagation in porous media." *J. Appl. Phys.*, 33(4), 1482–1498.
- Chen, T. W., Huang, L. H., and Song, C. H. (1997). "Dynamic response of poroelastic bed to nonlinear water waves." *J. Engrg. Mech.*, ASCE, 123(10), 1041–1049.
- Dean, R. G., and Dalrymple, R. A. (1991). *Water wave mechanics for engineers and scientists*, World Scientific, Singapore.
- Fenton, J. (1985). "A fifth-order Stokes theory for steady waves." *J. Wtrwy., Port, Coast., and Oc. Engrg.*, ASCE, 111(2), 216–234.
- Huang, L. H., and Chwang, A. T. (1990). "Trapping and absorption of sound waves. II: A sphere covered with a porous layer." *Wave Motion*, 12, 401–414.
- Huang, L. H., and Song, C. H. (1993). "Dynamic response of poroelastic bed to water waves." *J. Hydr. Engrg.*, ASCE, 119(9), 1003–1020.
- Liu, L. F. (1973). "Damping of water waves over porous bed." *J. Hydr. Div.*, ASCE, 99(12), 2263–2271.
- Madson, O. S. (1978). "Wave-induced pore pressure and effective stress in a porous bed." *Geotech.*, 28, 377–393.
- Mei, C. C. (1983). *The applied dynamics of ocean waves*, Wiley, New York.
- Mei, C. C., and Foda, M. A. (1981). "Wave-induced responses in a fluid filled poroelastic solid with a free surface—a boundary layer theory." *Geophys.*, 66, 597–631.
- Moshagen, H., and Torum, A. (1975). "Wave induced pressures in permeable seabeds." *J. Wtrwy., Harb. and Coast. Engrg. Div.*, ASCE, 101(1), 49–57.
- Putnam, J. A. (1949). "Loss of wave energy due to percolation in a permeable sea-bottom." *Trans. Am. Geophys. Union*, 30, 349–356.
- Reid, R. O., and Kajiura, K. (1957). "On the damping of gravity waves over a permeable sea bed." *Trans. Am. Geophys. Union*, 30, 662–666.
- Sleath, J. F. A. (1970). "Wave induced pressure in bed of sand." *J. Hydr. Div.*, ASCE, 96, 367–378.
- Yamamoto, T., Koning, H. L., Sellmeijer, H., and Hijum, E. V. (1978). "On the response of a porous-elastic bed to water waves." *J. Fluid Mech.*, Cambridge, U.K., 87, 193–206.

APPENDIX III. NOTATION

The following symbols are used in this paper:

- a = amplitude of incoming water wave;
- \mathbf{D}^* = displacement vector of fluid in porous medium;
- \mathbf{d}^* = displacement vector of solid skeleton;
- G, λ = Lamé's constants of elasticity;
- h = mean water depth of channel;
- $i = \sqrt{-1}$;
- K = bulk modulus of compressibility of fluid;
- k_j = wave numbers in porous medium, $j = 1, 2, 3$;
- k_p = specific permeability;
- k_0 = wave number of incoming water wave;
- n_0 = porosity;
- $P^{*(1)}$ = perturbed pressure in channel;
- $P^{*(2)}$ = perturbed pressure in porous medium;
- P_0 = perturbed pressure on bed;
- \mathbf{S}^* = normal tensor of fluid;

ε_1 = first expansion parameters of Stokes wave, $k_0 a$;
 ε_2 = second expansion parameters of Stokes wave, k_0/k_2 ;
 η^* = displacement of wave deviated from mean free surface;
 Λ, Π, Ψ = Mach numbers of two longitudinal waves and one transverse wave;
 μ = dynamic viscosity of fluid;
 ξ^* = displacement of wave deviated from mean channel-bed interface;

ρ_s = density of solid;
 ρ_0 = density of water;
 $\boldsymbol{\sigma}^*$ = solid stress tensor;
 $\boldsymbol{\tau}^*$ = effective solid stress tensor;
 $\Phi^{*(1)}$ = velocity potential of channel flow;
 $\Phi_j^{*(2)}$ = j th kind of displacement potentials of porous medium, $j = 1, 2, 3$; and
 ω, Ω = angular frequencies of water wave, in $O(\varepsilon_1)$ and $O(\varepsilon_1 \varepsilon_2)$, $\Omega = \omega$; in $O(\varepsilon_1^2)$, $\Omega = 2\omega$.