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Understanding Space-time Patterns of Groundwater System by Empirical Orthogonal Functions: a Case Study in the Choshui River Alluvial Fan, Taiwan

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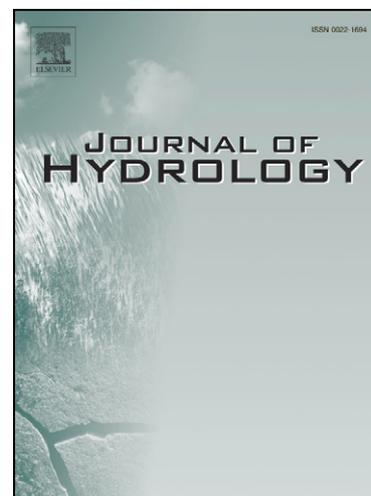
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1 Understanding Space-time Patterns of Groundwater System by Empirical
2 Orthogonal Functions: a Case Study in the Choshui River Alluvial Fan,
3 Taiwan

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5 **Abstract**

6 Natural or anthropogenic activities contribute to changes of groundwater levels in
7 space and time. Understanding the major and significant driving forces to changes in
8 space-time patterns of groundwater levels is essential to groundwater management.
9 This study analyzes monthly observations of piezometric heads from sixty-six wells
10 during 1997-2002 located in the Choshui River alluvial fan of Taiwan, where
11 groundwater has been the important local water resource for myriads of agricultural or
12 industrial demands. Following spatiotemporal estimations of piezometric heads by
13 Bayesian Maximum Entropy method (BME), this work performs rotated empirical
14 orthogonal function (REOF) analysis to decompose the obtained space-time heads
15 into a set of spatially distributed empirical orthogonal functions (EOFs) and their
16 associated uncorrelated time series. Results show that the leading EOFs represent the
17 most significant driving forces to spatiotemporal changes of groundwater levels in the
18 Choshui River aquifer. These include rainfall recharges from upstream Choshui and
19 Pei-Kang River, pumping activities from aquaculture usages in the coastal areas, as
20 well as water exchanges between surface and subsurface flow of Choshui River. In
21 summary, this study shows the strength of the REOF analysis which can effectively
22 provide integrative views of spatiotemporal changes of groundwater, gaining insights
23 of interactions between the groundwater system and other natural and human
24 activities.

25 **Keywords:** Empirical Orthogonal Function; Choshui River alluvial fan; Groundwater;
26 Bayesian Maximum Entropy; Space-time data analysis.

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27 **Introduction**

28 Groundwater has long been a reliable water source for a variety of uses such as
29 domestic, agricultural, and industrial uses (Yang and Yu, 2006; Chu and Chang,
30 2009). In Taiwan, comparing to the surface water supply, several major aquifers
31 provide the stable water resource, which accounts for over 30% of total amount of
32 water supply annually. However, the changes of groundwater level are the responses
33 of a complex interplay of a variety of natural and anthropogenic activities interacting
34 with the groundwater system. In addition, the prevalence of the heterogeneity in
35 subsurface environment and the lack of sufficient site characterization can
36 significantly hamper the understanding of space-time groundwater flow and transport
37 patterns, therefore, jeopardize the management of groundwater resources (Tartakovsky,
38 2007). Many studies about the changes of groundwater level often concentrate at the
39 changes induced by a specific driving force, such as natural and artificial recharges
40 from river or irrigation, pumping activities varying in space and time, and other
41 driving forces like earthquakes (Liu et al., 2004). Addition to extensive studies on the
42 impacts from specific activities by human or natural forces, it also requires the
43 systematic and integrative studies to obtain the macroscopic view of spatiotemporal
44 changes in hydraulic heads of the groundwater system of interest. For purposes of
45 groundwater management, it is essential to be knowledgeable about the major
46 ongoing underlying processes in space and time in the aquifer as well as the
47 magnitude of these processes that contribute to varying piezometric heads at specific
48 spatial and temporal locations.

49 In the groundwater monitoring investigations, the collected data may harbor
50 significant complex or extremely complicated variations in the observed values of
51 measurable characteristics of the groundwater level in time and space. Empirical

52 Orthogonal Function (EOF) analysis is an effective method to extract information
53 from large datasets in time and space domains (North, 1984; Weare and Nasstrom,
54 1984; Kim and Wu, 1999; Hannachi et al., 2007; Munoz et al., 2008). EOF analysis
55 conducts to the decomposition of the covariance kernel on the set of its
56 eigen-functions. EOF analysis reduces the dimensions of a space-time random data
57 fields into a smaller set of new spatial random fields which can be fairly accurate to
58 reconstruct the space-time variances of the original random fields (Hannachi et al.,
59 2007). Moreover, the purpose of the technique is to fit orthogonal functions to a set of
60 observed data, resulting in a reduction in the amount of data with minimal loss of
61 information while capturing the essential features (Munoz et al., 2008). Meteorology
62 has applied EOF analysis for decades to extract the most significant spatial signals of
63 atmospheric fields (Hannachi et al., 2007). Due to its advantage to obtain snapshots of
64 essential pure spatial and/or temporal patterns of a space-time dataset, many other
65 disciplines have recently applied EOF analysis in spatiotemporal analysis, such as
66 ozone distribution (Fiore et al., 2003), and ecological processes (Bejaoui et al., 2008).
67 For groundwater studies, EOF analysis was applied to extract significant temporal
68 signals from the Rhine Valley aquifer located in France and Germany (Longuevergne
69 et al., 2007), and was used to reduce the space-time variable to replace a large
70 groundwater numerical model by a comparable reduced model (McPhee and Yeh,
71 2008; Vermeulen et al., 2004). For environmental monitoring, Munoz et al. (2008)
72 considered Mid-Atlantic Stream Probabilistic Survey conducted from 1998 to 2002,
73 incorporated the spatio-temporal information in sampling designs, and illustrated how
74 to use the EOF model estimating at non-observed sites.

75 The study applied the EOF analysis to the case study in the Choshui River Fan
76 aquifer of Taiwan. The local farmers converted their crop lands into more profitable

77 aquaculture ponds, owing to abundant and low cost groundwater resources. Due to the
78 lack of an effective groundwater management policy in Taiwan, heavy groundwater
79 usage by aquaculture activities has notoriously caused local land subsidence,
80 sea-water intrusion, and aquifer salinization (Hsu, 1998; Liu et al., 2003; Liu et al.,
81 2006). Since 1992, the Water Resources Agency of Taiwan initiated a groundwater
82 monitoring network plan (GMNP) to systematically establish groundwater monitoring
83 wells with a spatial density of about $20\text{km}^2/\text{station}$ throughout major aquifers in
84 Taiwan (Hsu, 1998), to gather essential information including groundwater quality
85 and level, as well as hydrogeologic characteristics of the aquifers. The evenly
86 distributed space-time observations collected by the GMNP consist of a valuable
87 database containing comprehensive information about spatiotemporal variation of
88 groundwater quality and levels induced by a variety of physical and chemical
89 processes. The database provides essential information for a myriad of hydrogeologic
90 researches in areas covered by the GMNP, including the study area.

91 This study used EOF analysis to obtain the most significant spatially distributed
92 processes (i.e. EOFs) and their associated temporal variation from space-time
93 groundwater observations from the Choshui River alluvial fan. Before EOF analysis,
94 the current work performed spatiotemporal interpolation by the Bayesian Maximum
95 Entropy (BME) method to generate evenly distributed space-time estimations to
96 minimize potential systematic biases from sampling. During the analysis, EOF
97 rotation played an important role to extract the most informative signals from the
98 observations. This study then identified and interpreted the leading driving forces in
99 groundwater level variation.

101 **Materials and Method**

102 **Study area**

103 The Choshui River alluvial fan is located on the mid-western coast of Taiwan, and
104 covers the fertile plain area of 1800 km² including counties of Yun-Lin, Chang-Hua,
105 and northern Chia-Yi, as Figure 1 shows. Across the Choshui River, the largest river
106 in Taiwan, the alluvial plain is surrounded by natural geographical boundaries of the
107 Taiwan Strait to the west, the Central Mountain Ridge to the east, the Wu River to the
108 north, and the Pei-Kang River as its southern border. Annual rainfall in this area is
109 around 2460 mm and 78 percent of precipitation occurs from May to October, i.e.
110 plum rain and typhoon seasons. The annual runoff in the Choshui River is about 6.08
111 billion tons (Chen and Lee, 2003). Because of insufficient surface water supply in the
112 alluvial fan, residents extract groundwater to supplement their demands irrigation,
113 aquaculture, and household, particularly in dry seasons. Among them, groundwater is
114 the major clean water supply for aquaculture ponds and therefore residents illegally
115 extract a great amount of water from aquifers into aquaculture ponds. The overdraft of
116 groundwater in agriculture and fish cultivation is causing serious land subsidence in
117 coastal areas (Yang and Yu, 2006).

118 The Choshui River alluvial fan is partitioned primarily into proximal-fan, mid-fan and
119 distal-fan areas, according to their distinct hydrological formations. Figure 2 shows
120 the conceptual hydro-geological profile in the Choshui River alluvial fan. The
121 hydrogeological formation consists of three major aquifers, i.e. aquifer I, II, and III
122 numbered from the ground surface level, and separated by the aquitards, which are
123 low permeable with fine sediment, ranging from clay to fine sand. Considering
124 hydrogeological formation, the proximal-fan is the major recharge area of the aquifer

125 (Jang et al., 2008; Jang and Liu, 2004). The aquitards located in the distal-fan and
 126 mid-fan areas gradually diminish in thickness toward the east. Moreover, Aquifer II is
 127 the major aquifer of the Choshui alluvial plain because of its large spatial extent and
 128 acceptable depth for groundwater retrieval (Liu et al., 2004). Data derived from
 129 pumping tests indicate that the observed hydraulic conductivity fields ranges from
 130 10^{-3} – 10^{-5} m/s, and decreases from the proximal fan to the distal fan (Hsu, 1998; Jang
 131 et al., 2008; Jang and Liu, 2004). Transmissivity ranges from 0.04–4.19 m²/min. The
 132 storage coefficient is about 0.1 for the unconfined aquifer and ranges from 10^{-3} – 10^{-4}
 133 for the confined aquifer (Hsu, 1998). In this study, the dataset includes pizeometric
 134 head observations of aquifer II obtained from sixty-six monitoring wells, evenly
 135 distributed over the entire Choshui River alluvial fan. The study recorded the
 136 observations monthly during the period from July 1997 to December 2001.

137

138 **Method**

139 The aim of EOF analysis is to decompose a continuous space-time random field

140 $X(s,t)$ into the additive space-time multiplication form as follows

$$141 \quad X(s,t) = \sum_{k=1}^M c_k(t)u_k(s) \quad (1)$$

142 where the vector (s,t) denotes the space-time location at time t and spatial position

143 s . M is the number of modes in orthogonal space-time random fields, i.e.

144 $c_k(t)u_k(s)$. The modes are formulated as an optimal set of orthogonal spatial functions

145 $(u_k(s))$, i.e. EOFs, and their associated expansion functions of time $(c_k(t))$, i.e., the

146 projection of $X(s,t)$ on $u_k(s)$, also called EOF expansion coefficients (ECs). The

147 concept of EOF analysis is essentially conventional principal component analysis
 148 (PCA), which generates a smaller set of new random variables. The major leading
 149 EOFs can usually explain the fairly amount of the observed variances of the original
 150 space-time dataset, e.g. in this study, five EOFs can explain over 80% of the variances
 151 of space-time groundwater head data as shown below. To consider the geometrical
 152 relationship among the space-time dataset, not common in most PCA applications,
 153 this work first interpolates space-time observations into regularly spaced grids over
 154 the entire space-time domain. This mitigates data clustering effects, which can
 155 contribute to excess variances of clustering locations, therefore distorting EOF
 156 analysis results (Buell, 1971; Buell, 1978; Karl et al., 1982). This study uses the
 157 Bayesian maximum entropy method (BME) to estimate the spatiotemporal
 158 distribution of piezometric heads by accounting for spatiotemporal dependence, i.e.
 159 covariance, as well as for observations considered as hard data in this case. For a more
 160 detailed description of the BME method, the reader can refer to the literature
 161 (Christakos, 2000; Christakos et al., 2002).

162 In EOF analysis, the head covariance over spatial domain finds the uncorrelated
 163 spatial functions such that $Cu = \lambda^2 u$, where C is the covariance among the gridded
 164 data in space, $u = (u_1, \dots, u_p)^T$ is the matrix that composes the eigenvectors u_k
 165 corresponding to eigenvalues (λ_k) , and p is the number of space locations. Without
 166 generality loss, the spatial (temporal) covariance (C) can be expressed as $C = \frac{1}{n} XX^T$
 167 where X is a $p \times n$ matrix containing space-time BME estimations of piezometric
 168 heads with the number of observed time (n). The amount of observed head variance
 169 explained by the eigenvector (u_k) is the value of its associated eigenvalue (λ_k) .

170 In practice, the singular value decomposition (SVD) method is used for EOF analysis
 171 (Hannachi et al., 2007). A $p \times n$ matrix of space-time head estimations $X(s,t)$ can
 172 decompose as

$$173 \quad X = U\Lambda A^T \quad (2)$$

174 where U and A respectively $p \times M$ and $n \times M$ are the unitary matrix, i.e.
 175 $U^T U = A^T A = I$, in which the columns $u_k(s)$ are essentially EOFs as the spatial
 176 orthonormal basis of the space-time data matrix. The diagonal matrix (Λ) with
 177 elements of $\lambda_1 \geq \lambda_2 \geq \dots \geq \lambda_r$ are singular values of the matrix of $X(s,t)$. Therefore, the
 178 projections of EOFs (c_k) are expressed as $c_k(t) = \lambda_k a_k(t)$. The space-time
 179 decomposition of Eq. (1) by EOF analysis can be rewritten as

$$180 \quad X(s,t) = \sum_{k=1}^M \lambda_k a_k(t) u_k(s) \quad (3)$$

181 One of the major challenges for EOF analysis is to interpret the estimated EOFs and
 182 their associated projections which are orthogonal to each other but may not be
 183 physically meaningful. The rotation of EOF patterns (REOF) can be one of the most
 184 common approaches to overcome the interpretation issue (Hannachi et al., 2007). The
 185 rotation concept systematically alters the original EOF structure based upon some
 186 criterion, such as maximizing the explained variances of leading EOFs. Studies of
 187 multivariate statistical analysis, e.g. factor analysis (Anderson, 2003) have proposed
 188 and widely applied a variety of rotation algorithms. Among them, the Varimax method
 189 is the most well-known and used rotation technique, by which an orthogonal matrix is
 190 applied to EOF rotation to simplify the EOF structure, pushing the loading
 191 coefficients of EOFs to either zeros or ± 1 (Kaiser, 1958). Determining the number
 192 of modes, M , for space-time decomposition is also a major issue in EOF analysis.

193 For reducing dimension, the value of M is always chosen to be much less than the
194 numbers of space-time dimensions of observations, i.e. n and p . However, the
195 general outcome of REOF analysis depends on the selection of M which can
196 complicate understanding of the underlying leading physical patterns of the
197 observations. To obtain the invariant leading REOFs, EOFs should be re-scaled
198 according to their associated eigenvalues before rotation. Rotating rescaled EOFs
199 generates invariant leading REOFs due to relatively little contributions from the
200 scaled EOFs of smaller eigenvalues (Hannachi et al., 2007). The current study
201 considers the varying piezometric head of the aquifer as the linear superposition of
202 several contributions from independent natural or anthropogenic processes
203 decomposed by REOF analysis.

204

205 **Results**

206 *Spatiotemporal distribution of piezometric heads*

207 This research predicted monthly spatial distributions of piezometric heads of aquifer
208 II in the Choshui River alluvial fan by the BME method, accounting for the
209 spatiotemporal trend and covariance among the heads. Figure 3 shows the piezometric
210 heads results of two selected months and the triangles represent the monitoring wells.
211 The highest piezometric head is at the proximal-fan of the Choshui River alluvial fan
212 and the lowest is close to the southern coastal area i.e. Yi-Wu (Figure 3 (a) and (b)).
213 The hydraulic gradient from east to west is caused by topography changes in the
214 presented area significantly. The distribution of piezometric heads changes slightly
215 from month to month, primarily from obvious seasonal precipitation over the study
216 area, i.e. central Taiwan. The wet and dry seasons are from May to October and from

217 November to April, respectively. Piezometric heads over the Choshui River alluvial
218 fan in March (Figure 3 (a)) and October (Figure 3 (b)) in 2001 represent the general
219 spatial distribution of the groundwater level during the dry and wet seasons,
220 respectively. The both figures show a slight difference in the piezometric heads
221 between the two seasons, especially in the coastal areas.

222 *The EOFs in spatial domain and Time series of the ECs*

223 Figure 4 shows total variance among the observed head data primarily explained by
224 the leading EOFs. Among them, the first five EOFs explain about 80 percent of the
225 observed spatiotemporal changes of heads with their contributions of 47.9%, 10.8%,
226 9.8%, and 6.9% and 4.7%, respectively. Figure 5 shows spatial distributions of the
227 first five EOFs. Each EOF has its distinct spatial pattern, which is generally localized.
228 Figure 6 shows the associated ECs (shown as the black line) during the study period
229 of the five EOFs. Equation (1) shows that jointly considering ECs and EOFs reveals
230 positive or negative contributions from each of the EOFs to piezometric head changes
231 in space and time. This study uses the bright areas (hotspot) of EOFs to represent
232 positive contributions to piezometric head changes. In EOF1, the brightness hotspot is
233 located upstream to the Choshui River, primarily in the Gu-Keng and Dou-Liu
234 townships, shown in Figure 5 (a). The EOF3 also shows a similar spatial pattern
235 where the brightness area is located upstream to the Pei-Kang River, shown in Figure
236 5 (c). As mentioned in previous studies (Jang et al., 2008; Jang and Liu, 2004), the
237 proximal-fan is a major recharge region for aquifers, due to its hydro-geological
238 formation being primarily composed of gravel and sand. Compared to rainfall
239 observations at the Da-Pu station located upstream to the Choshui River, Figure 6 (a)
240 shows that EC1 temporal variation highly associates with the hydrologic cycle in the
241 area, yet with about two or three months delay, the approximate time required for

242 rainfall to percolate in the aquifer. The temporal pattern of EC1 implies that upstream
243 recharge from the Choshui River is the leading driving force causing spatiotemporal
244 changes of the aquifer. This study also observes a similar rainfall recharging pattern in
245 EC3 in which the trend varies closely to rainfall measurements at Dou-Nan, located
246 upstream to the Pei-Kang River, i.e. the brightness area of EOF3.

247 The EOF2 hotspot is near Yi-Wu and King-Hu, among the greatest land subsidence
248 locations (TPWCB, 1996; TPWCB, 1997), shown in Figure 5 (b). Because of heavy
249 aquaculture and irrigation demands over the entire township, illegal over-pumping
250 resulted in groundwater level decline (Akudago et al., 2009), therefore consolidating
251 soil layers (Liu et al., 2004). Cumulative land subsidence amounts from 1976 to 2000
252 obtained by a leveling survey at Yi-Wu and King-Hu (within the Ko-Hu township)
253 were 195 and 188 cm, respectively. The temporal pattern of EC2 closely corresponds
254 to the variation of measured piezometric heads at a monitoring well close to Yi-Wu,
255 shown in Figure 6 (b). The piezometric head falls during spring and summer, and
256 arises during the other seasons. During the high season of water usage, particularly
257 from March to July, local farmers extract groundwater for irrigation, fish cultivation
258 and household demands and result in the seasonal drawdown of groundwater levels.
259 Figure 6(e) also shows the EC5 time series and piezometric heads in the Shi-Kong
260 gauge. The EC5 increases with time. In fact, the groundwater level has begun to
261 rebound and the subsidence rate in Shi-Kong has declined (Liu et al., 2004). The
262 government has not allowed intensive groundwater use due to the industrial
263 development in the area since 1998. Moreover, the EOF4 hotspot is in the Choshui
264 River, shown in Figure 5(d). Figure 6 (d) shows the streamflow during the study
265 periods in the Chang-Yun Bridge gauge and the EC4 varies with the streamflow. The
266 EOF4 driving force is the exchange between the Choshui River and groundwater.

267 Moreover, groundwater and surface water are not isolated components of the
268 hydrologic system, and interactions exist between ground water and surface water.
269 During flooding, the river recharges the aquifer. During the dry season, groundwater
270 flux drains into the stream, leading to increased stream flows (Sophocleous, 2002).

271

272 **Discussion**

273 Groundwater studies of the Choshui River alluvial fan have primarily focused on
274 issues driven by anthropogenic activities, such as land subsidence and its associated
275 impacts (Liu et al., 2004). An integrative study performed on the aquifers to identify
276 the major underlying processes of the groundwater system would be more helpful for
277 groundwater management. In this study, REOF analysis shows its effectiveness to
278 reveal, not only the leading driving forces of groundwater level changes in space and
279 time, but their interactions with the aquifer. The study by (Longuevergne et al., 2007),
280 also shows that EOF analysis reveals the primary characteristics at the Rhine Valley
281 aquifer (France and German). Both the study (Longuevergne et al., 2007) and our
282 study require an extensive groundwater monitoring network for the aquifer, i.e.
283 ninety-five and sixty-six monitoring wells for the Rhine Valley and Choshui aquifers,
284 respectively. Contrasted to the Rhine Valley study, our study performs spatiotemporal
285 interpolation of piezometric heads before EOF analysis to reduce the effects from
286 uneven spatial distribution of monitoring wells (Karl et al., 1982; Wikle and Cressie,
287 1999). The current study also shows that rotating EOFs effectively increases EOF
288 interpretability by generating more spatially localized and stable spatial patterns of
289 leading EOFs.

290 As shown in previous studies (Chen and Lee, 2003; Jang and Liu, 2004), recharges

291 play an important role in Choshui river aquifer sources, especially in the proximal-fan
292 area, i.e. upstream to the Choshui and Pei-Kang rivers. In the proximal-fan area, the
293 logarithm of hydraulic conductivities are generally higher, about 4-5 $\ln(\text{m/day})$, than
294 those in rest of the aquifer, about 1-3 $\ln(\text{m/day})$ (Jang and Liu, 2004). One of the most
295 valuable features of REOF analysis to groundwater analysis is its ability to clearly
296 identify primary recharge areas for the aquifer. Identifying recharge areas is an
297 important step towards protecting regional groundwater resources (Braun et al., 2003).
298 The inappropriate use of these areas increases the risk of groundwater contamination.
299 Moreover, identifying the recharge source could be useful for managing aquifers to
300 meet increasing demand, and also help address environmental issues on effects of
301 water level decline (Acheampong and Hess, 2000). In this study, time series EC1 and
302 EC3 highly associate with rainfall measurements of local weather stations. The
303 comparison between temporal variations of ECs and rainfalls, and percolation time for
304 groundwater recharge depends on several hydrogeological factors, including hydraulic
305 conductivity and depth of the groundwater table (Gau et al., 2006). The amount of
306 recharges closely relates to the amounts of local rainfalls and stream flows.
307 Quantifying groundwater recharge is typically difficult because direct recharge input
308 to the water table is not easily measured, especially when the water table is several
309 meters below the land surface in an aquifer (Gburek and Folmar, 1999), due to the
310 absence of effective instrumentation. By different techniques, the estimations of
311 annual groundwater recharge in the mountain region of Choshui aquifer range from
312 3.1 to 3.5 billion tons (Chen and Lee, 2003; Gau and Liu, 2000). In the proximal fan
313 of the Choshui aquifer, percolation time for rainfall to reach the groundwater table is
314 about two to three months, similar to results of the aquifer study in the central coastal
315 plain (Israel) (Rimon et al., 2007).

316 The study also identifies exchanges between the Choshui River flow and ground
317 water as the major contributing factor to changes in ground water level. The primary
318 contribution of the Choshui river flow is along the Choshui River, as expected.
319 Particularly, the most sensitive areas for river recharge are located upstream to the
320 Choshui River and the Pei-Kang River. The results also reveal interactions between
321 surface and subsurface water of the Choshui River. Changes in EC4 are much
322 smoother compared to streamflow changes. Moreover, temporal variation of flow rate
323 generally fluctuates significantly in response to rainfall. The changes in EC4 better
324 reflect the base flow temporal pattern of the Choshui River, which reacts slower than
325 the runoff to local rainfall changes. As a result, EC4 shares temporal characteristics
326 similar to EC1; however, with different magnitudes, i.e. EC1 is directly associated
327 with seasonal rainfall and EC4 is more connected to the flow pattern of Choshui River
328 which is closely related to rainfall.

329 As expected, this study identified pumping at several places in the coastal area, as
330 among the major contributing processes to piezometric head changes. In the area, the
331 soil consists mostly of clay and fine sand; the strength and the permeability of this soil
332 are relatively low (Liu et al., 2004). Furthermore, the Choshui River alluvial fan
333 includes the major aquaculture towns in Taiwan, and therefore extensive groundwater
334 demands are expected, because of insufficient surface water supply in these areas.
335 Illegal overpumping of groundwater has been prevalent in almost the entire coastal
336 area counties of Chang-Hua and Yun-Lin since 1950 due to a lack of ground water
337 management. The accumulated land subsidence due to unmanaged pumping activities
338 ranges from 50cm to 200cm along the coastline of the Choshui River alluvial fan (Liu
339 et al., 2004). This study identified two hotspots where pumping activities were still
340 active during the study period of 1997-2002, i.e. the piezometric heads still changed

341 significantly. Among them, fortunately, the EC5 shows that unmanaged pumping
342 seems to be under control and the groundwater level in Shi-Kang gauge has started to
343 rebound consistently since mid-1998 at the hotspot area of EOF5. The Ko-Hu
344 township identified by EOF2 shows that the regular seasonal pattern of hydraulic
345 heads at its lowest time occurred in spring and summer every year during the study
346 period. The seasonal pattern is closely associated to the water demands for local
347 aquaculture ponds (Yang and Yu, 2006).

348

349 **Conclusion**

350 This study presented a macroscopic and integrative approach to investigate the
351 spatiotemporal changes of a groundwater system by REOF analysis. We analyzed the
352 monthly records of groundwater levels from 1997 to 2002 for sixty-six monitoring
353 wells operated by a water resources agency in Taiwan. This study shows that REOF
354 analysis can effectively capture stable and localized features, and gain easy
355 interpretation of EOFs. The current study identified five underlying processes as
356 major contributors to changing groundwater levels in the aquifers of Choshui River
357 alluvial fan, including recharges from rainfalls, stream flow and groundwater usage in
358 the coastal areas. These five leading EOFs drive the system changes, amounting to
359 about 80 percent of global variance for the entire groundwater system. More
360 specifically, the sensitive recharge areas are located upstream to the Choshui River
361 and the Pei-Kang River. This finding suggests a required groundwater management
362 policy in these places to ensure avoiding any potential contamination. Though land
363 subsidence is prevalent along the coastline of the Choshui River alluvial fan, the
364 locations with the most significant groundwater level changes are near the coastal area,

365 i.e. townships of Ko-Hu (Yun-Lin county), and Da-Cheng (Chang-Hua county) in our
366 analysis.

367 This study shows the REOF analysis can effectively reveal the underlying space-time
368 processes of groundwater system. The REOF analysis results provide the integrative
369 view of interests in groundwater system and insights of major contributing factors to
370 the groundwater level changes in space and time, which are the essential information
371 for the effective management of a groundwater system.

372

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480 Figure Captions

481 **Figure 1** Geographical location of the Choshui River alluvial fan in Taiwan

482 **Figure 2** Conceptual hydro-geological profile of the Choshui River alluvial fan

483 **Figure 3** The piezometric head (in meter) maps using BME on (a) March, 2001, and
484 (b) October, 2001

485 **Figure 4** Variance percentage of rank of EOFs

486 **Figure 5** The first five EOF interpolations (Unit: m)

487 **Figure 6** Time series of the ECs (black line) and the hydrologic components: (a)

488 rainfall distribution in Da-Pu (b) piezometric head distribution in Yi-Wu (c) rainfall

489 distribution in Dou-Nan (d) streamflow distribution in Chang-Yun Bridge (e)

490 piezometric head distribution in Shi-Kang

491

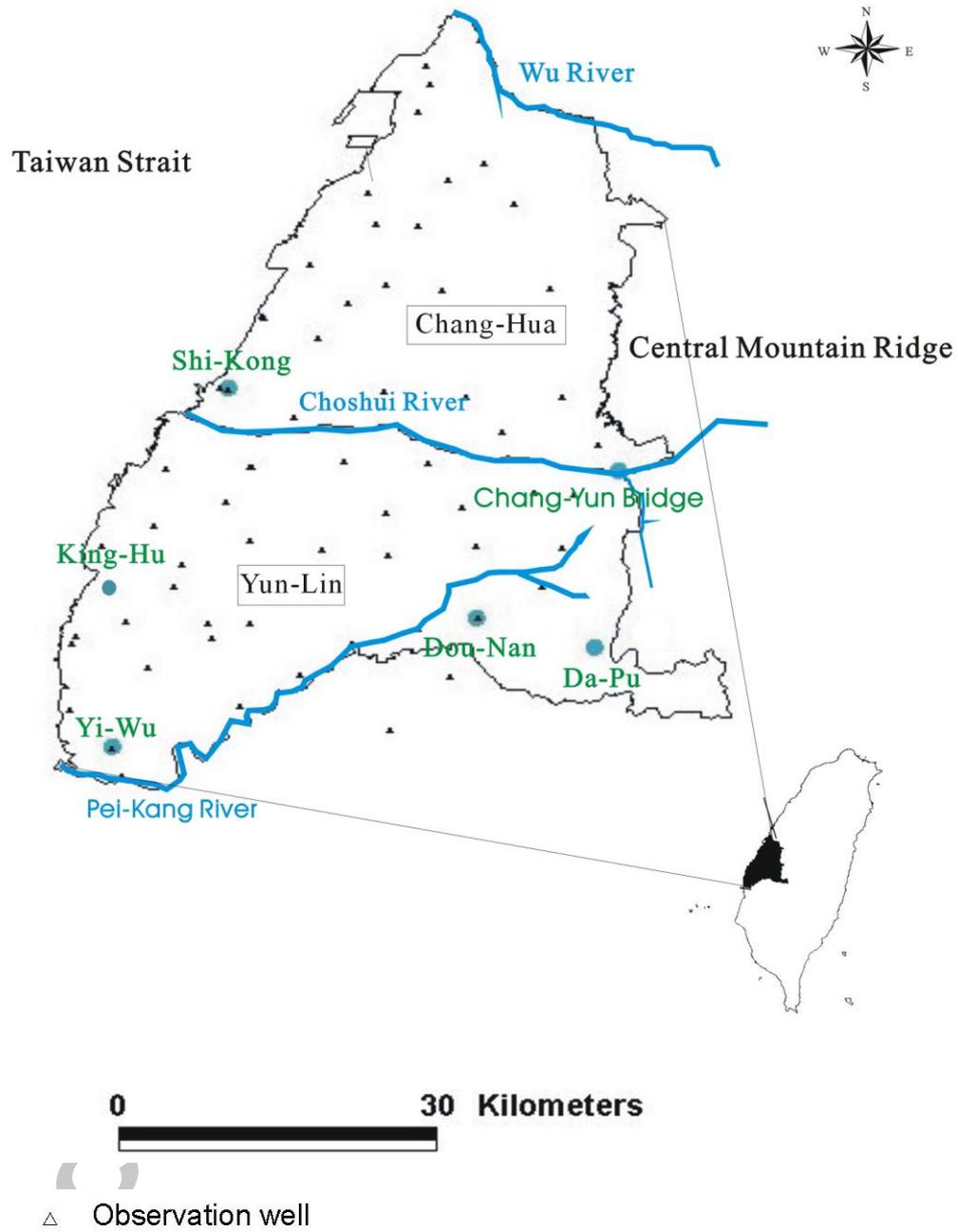


Figure 1 Geographical location of the Choshui River alluvial fan in Taiwan

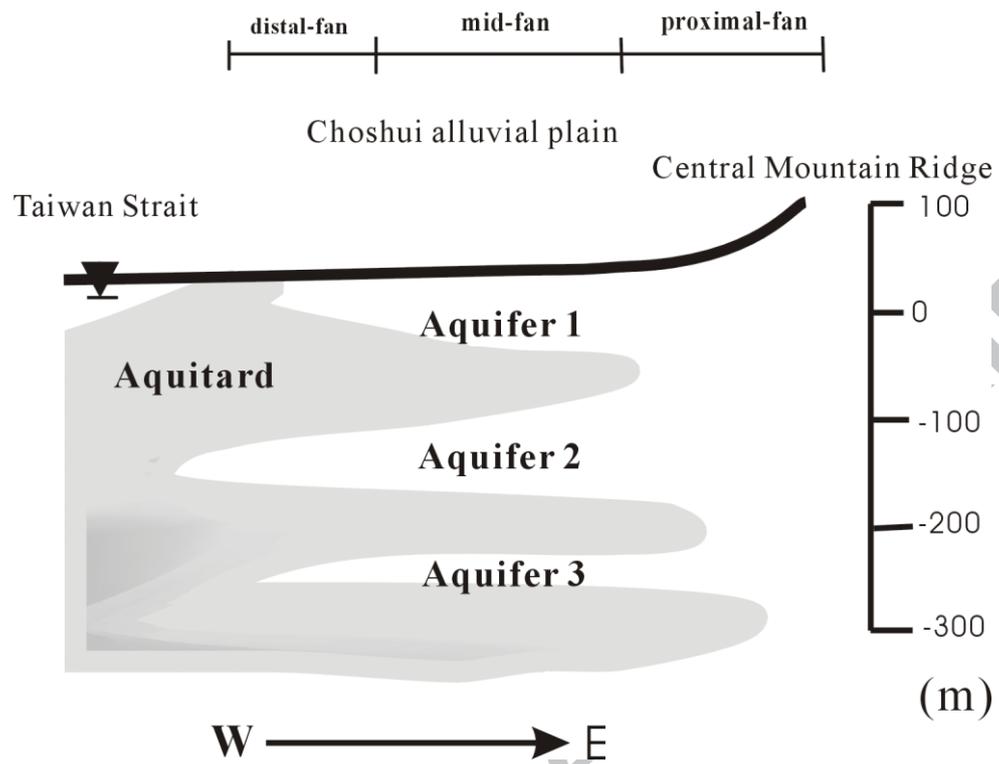
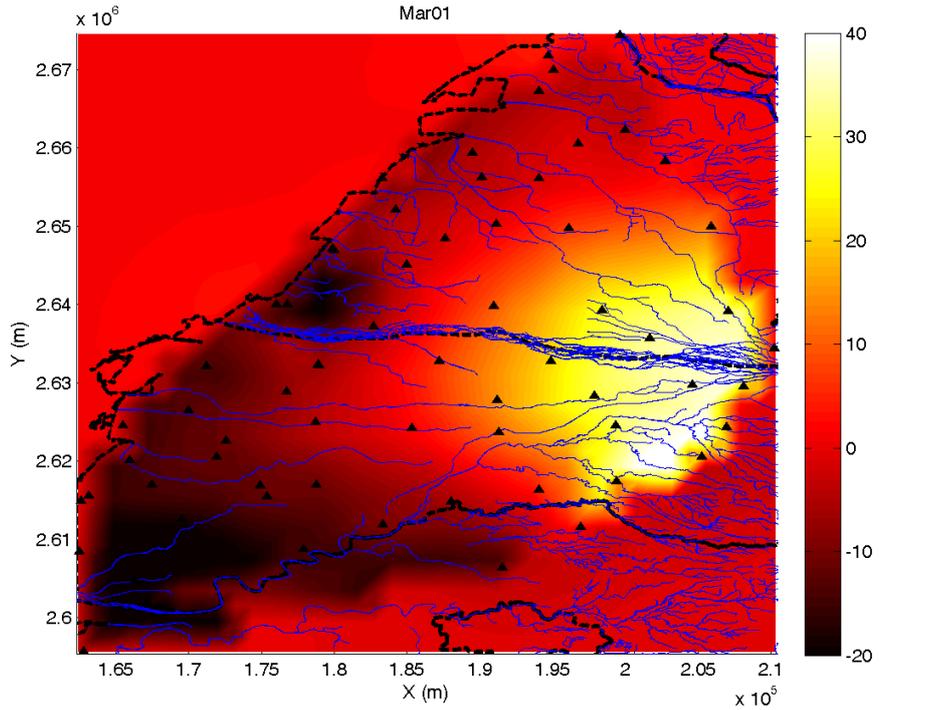
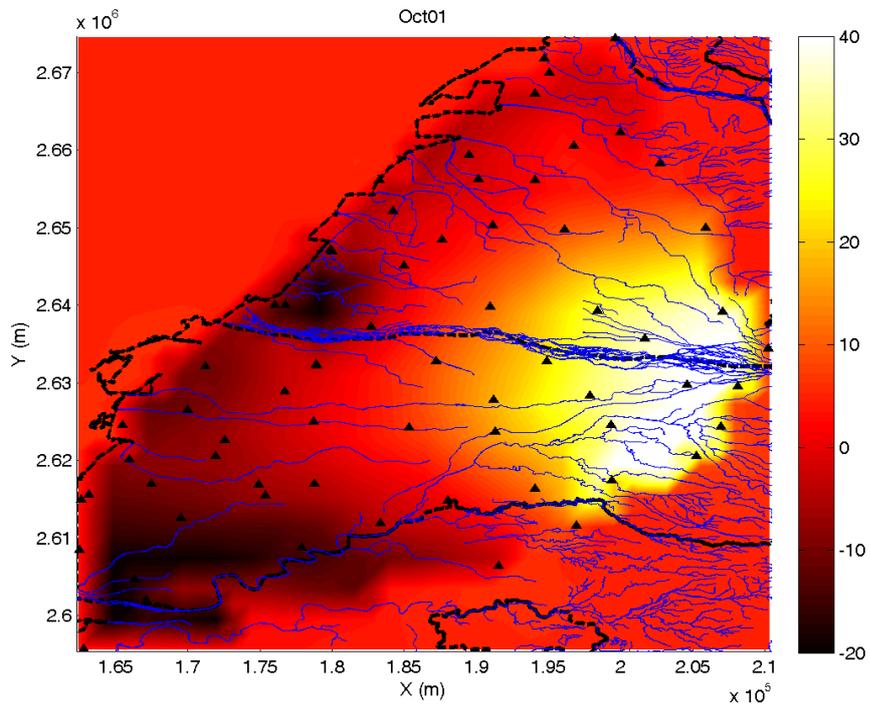


Figure 2 Conceptual hydro-geological profile of the Choshui River alluvial fan



(a)



(b)

Figure 3 The piezometric head (in meter) maps using BME on (a) March, 2001, and (b) October, 2001

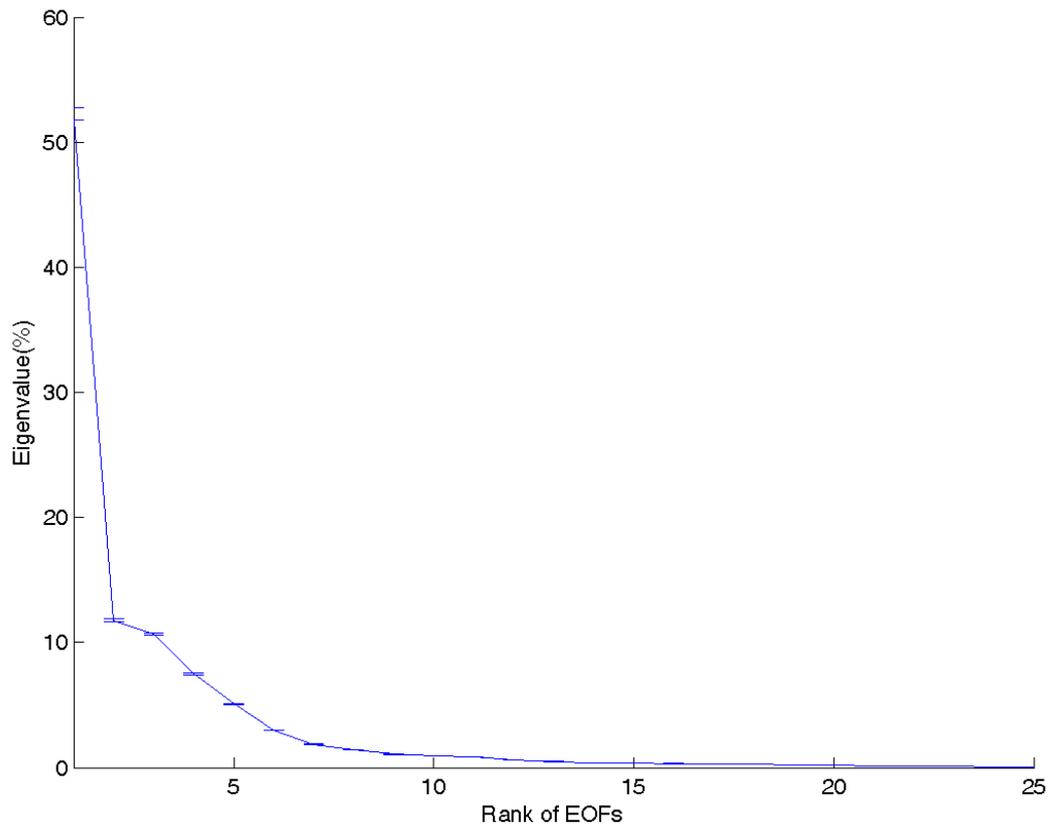


Figure 4 Variance percentage of rank of EOFs

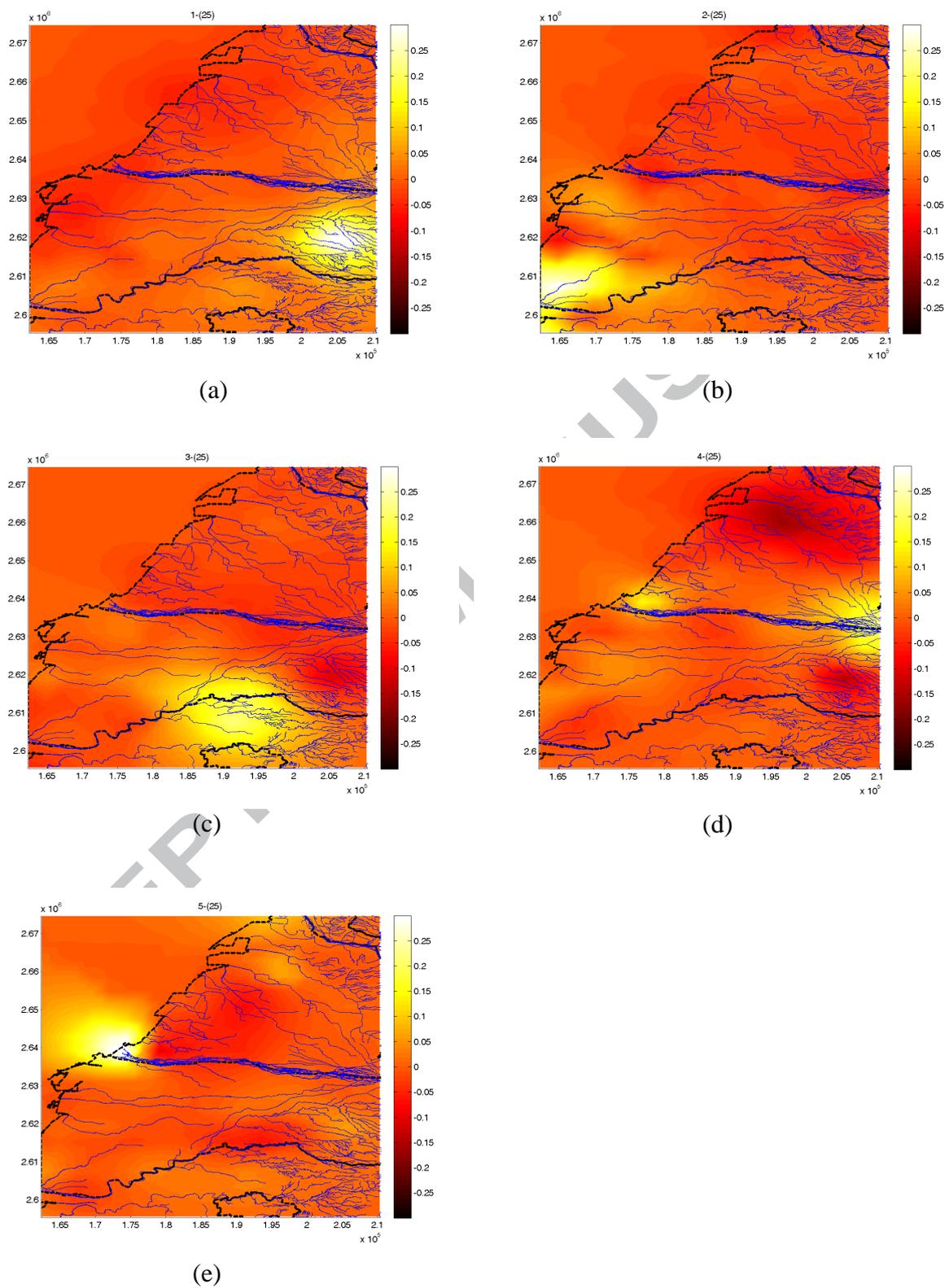


Figure 5 The first five EOF interpolations (Unit: m)

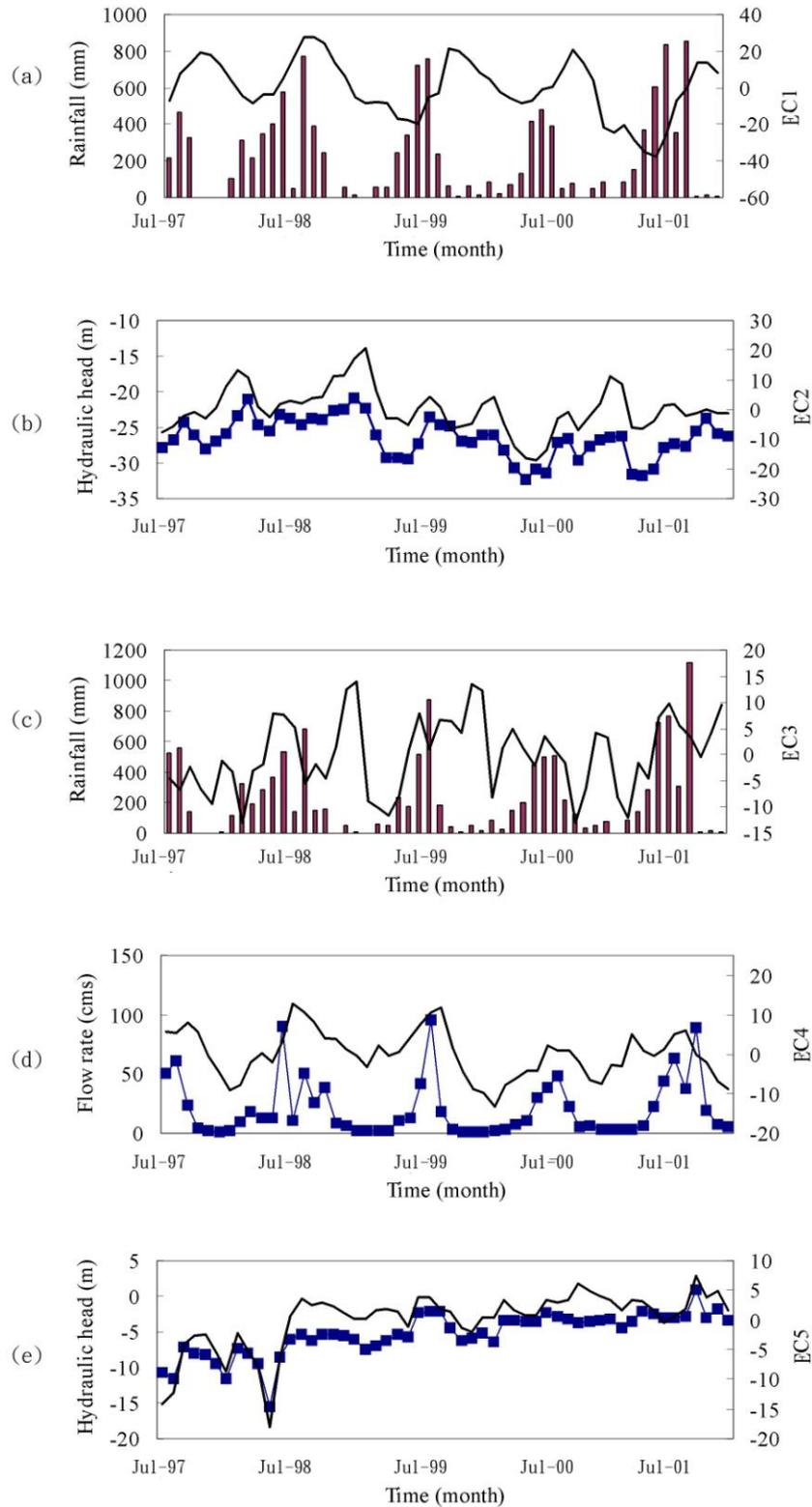


Figure 6 Time series of the ECs (black line) and the hydrologic components: (a) rainfall distribution in Da-Pu (b) piezometric head distribution in Yi-Wu (c) rainfall distribution in Dou-Nan (d) streamflow distribution in Chang-Yun Bridge (e) piezometric head distribution in Shi-Kang